

# The parameterization of subgrid scale orography in LMDZ

Mountains influence the dynamics of the atmosphere at different length scales:

- They force **gravity waves** that take angular momentum from the earth and transport it through the atmosphere over long distances
- **At large-scales** mountains contribute to the steady planetary wave, to the storm tracks, to the low-frequency variability
- F is the force exerted by an obstacle on a fluid
- **The drag** is the component of the force F that decelerates the fluid because it is opposite to the wind
- **The lift** is the component of the force that modifies the direction of the flow but does not decelerate it.

# Outline

- Drag controlled by gravity waves
- Lift and forcing of the steady planetary waves
- Why are gravity waves important for stratospheric circulation?
- Why are planetary waves important for stratospheric circulation?

# 1. drag\_noro : drag controlled by gravity waves

The Lott and Miller (1997) scheme treats the Subgrid Scale Dynamics controlled by the Gravity Waves

Non-dimensional height of the mountain:  $H_n = NH/U$   
H is the maximum height of the obstacle

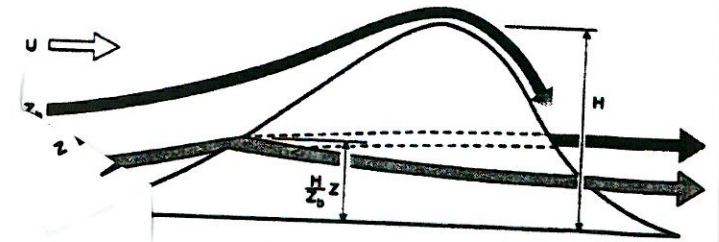
- At small  $H_n$  all the flow goes over the mountain, gravity waves are forced by the vertical motion of the fluid

Tau the surface stress due to the gravity wave :

$$\tau = \rho b G B(\gamma) N U H^2$$

- At large  $H_n$  the vertical motion of the fluid is limited and part of the low-level flow goes around the mountain for  $z < z_b$ .

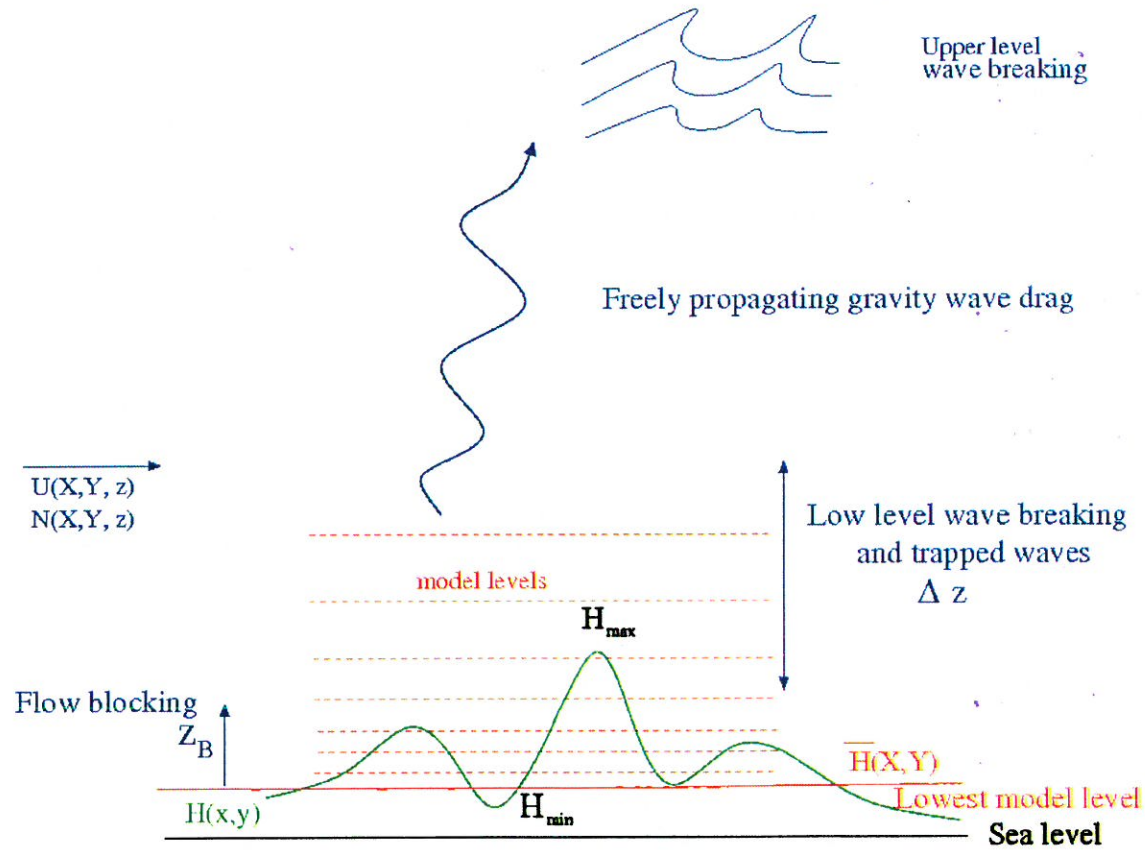
The surface stress is then :  $\tau = \rho C_d z_b b U^2$



# drag\_noro

The scheme depends on 4 parameters  $C_d$ ,  $G$ ,  $R_{ic}$  and  $z_b$ .

- $C_d$  and  $G$  control the amplitude of the blocked-flow drag and of the gravity-wave drag.
- $R_{ic}$  and  $z_b$  control the vertical distribution of these drags.



## 2. lift\_noro: The component of the force that modifies the direction of the flow

Pressure force acting on the mountain to the left due to the background pressure gradient associated with the mean flow:

$F = MV\rho U f$  ;  $MV$  is the mountain volume

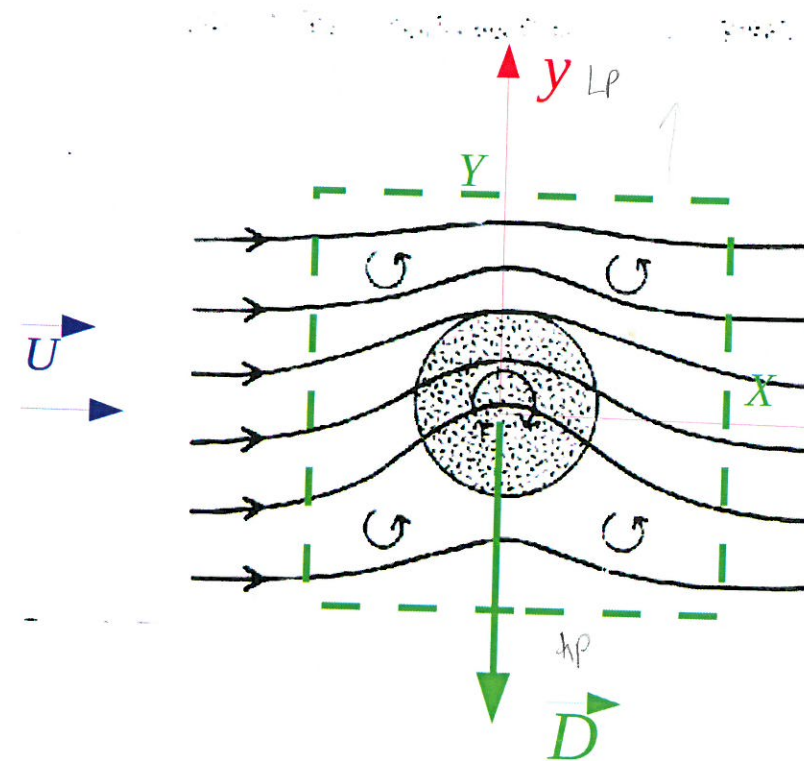
The lifting of the air over the mountain is balanced by the force acting on the mountain and then results in a push to the right (*Smith 1979*).

Between the narrow ridges of a mountain air can be blocked and separated from the large-scale flow.

A region of complex terrain acts as if it has a height larger than the actual height.

A solution could be to increase the sizes of the mountains

In lift-noro the mean orography is kept and the missing forces are applied.



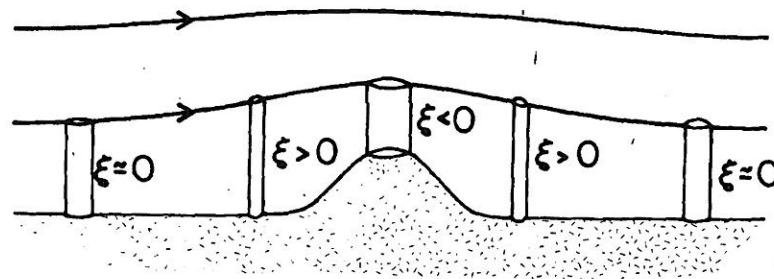


Conservation of potential vorticity for an adiabatic motion:

$$PV = \frac{1}{\rho} (f + \xi_r) \frac{d\theta}{dz}$$

If the motion to the South or North is large enough, one can't consider that  $f$  is constant.

**The mountain triggers a steady Rossby wave.**



### **3. Why are gravity waves important for the middle atmosphere circulation?**

The sources of gravity waves are orography, but also convective and frontal systems.

Gravity waves can propagate vertically and break.

The moment flux deposition due to gravity wave breaking is parameterized (for example in `hines_gwd`).

This breaking affects the mean circulation in the stratosphere (Quasi-Biennial Oscillation) and the mesosphere (changes in the zonal wind).

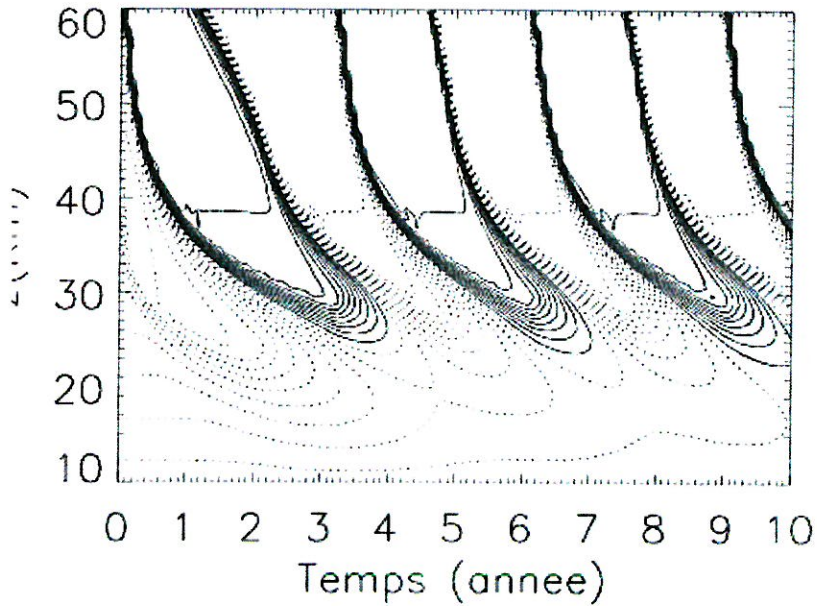
⇒ Quasi-Biennial Oscillation

Altitude de déferlement des ondes de gravité :-

$$z = 2H \ln \left( \frac{|w - ku|}{|m| w_0} \right)$$

- $\hat{\omega} > 0$  Accélère le vent moyen
- $\hat{\omega} < 0$  Freine le vent moyen

$\bar{u}$  (m/s)



①  $w > 0 \hat{\omega} > 0 \quad c_p > 0$   
 $u < 0 \Rightarrow |w - ku|$  grand  
 Pénétration de l'onde à haute altitude

④

accélération du vent  
déferlement

a) Départ  $U(z)$   $z$   $U'$   $-c$   $0$   $+c$

b)  $U(z)$   $z$   $U'$   $-c$   $0$   $+c$

$w < 0 \hat{\omega} < 0$   
 $c_p < 0$   
 $u < 0$   
 ⇒  $|w - ku|$  faible  
 Déferlement à basse altitude  
 Freine  
 → Anomalie vent d'Est

c) ⑤  $U(z)$   $z$   $U'$   $-c$   $0$   $+c$

l'anomalie de vent d'Est se déplace vers le bas

d)  $U(z)$   $z$   $U'$   $-c$   $0$   $+c$

Viscosité

e)  $U(z)$   $z$   $U'$   $-c$   $0$   $+c$

$u > 0 \quad c_p > 0$   
 $w > 0 \hat{\omega} > 0$   
 ⇒  $|w - ku|$  petit  
 Accélère le vent

f)  $U(z)$   $z$   $U'$   $-c$   $0$   $+c$

$u > 0 \quad w > 0 \hat{\omega} > 0$   
 $c_p > 0$   
 ⇒  $|w - ku|$  petit  
 Déferlement à basse altitude  
 Accélère le vent

le vent se déplace vers le bas

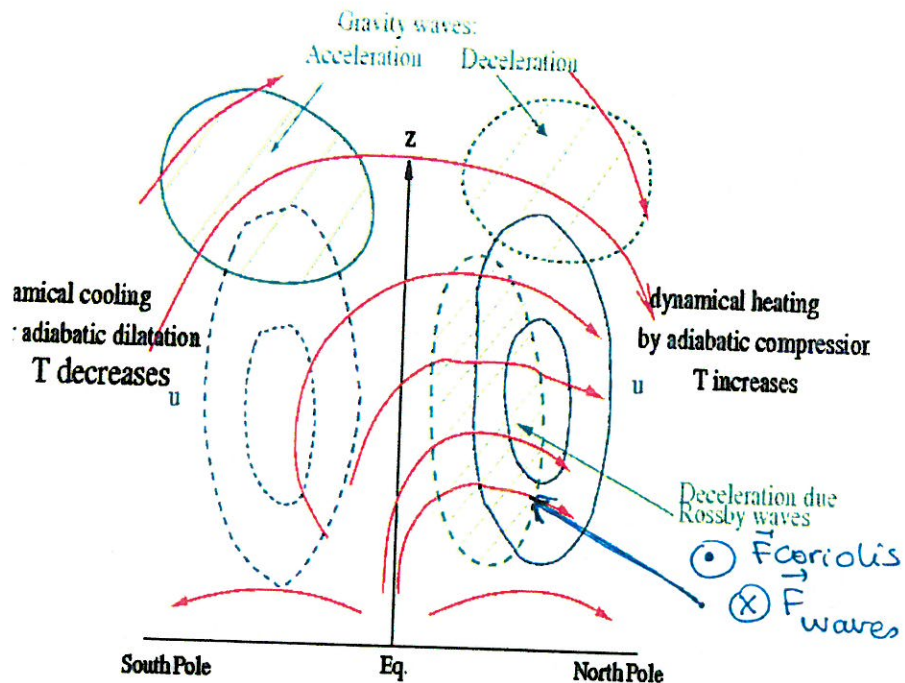
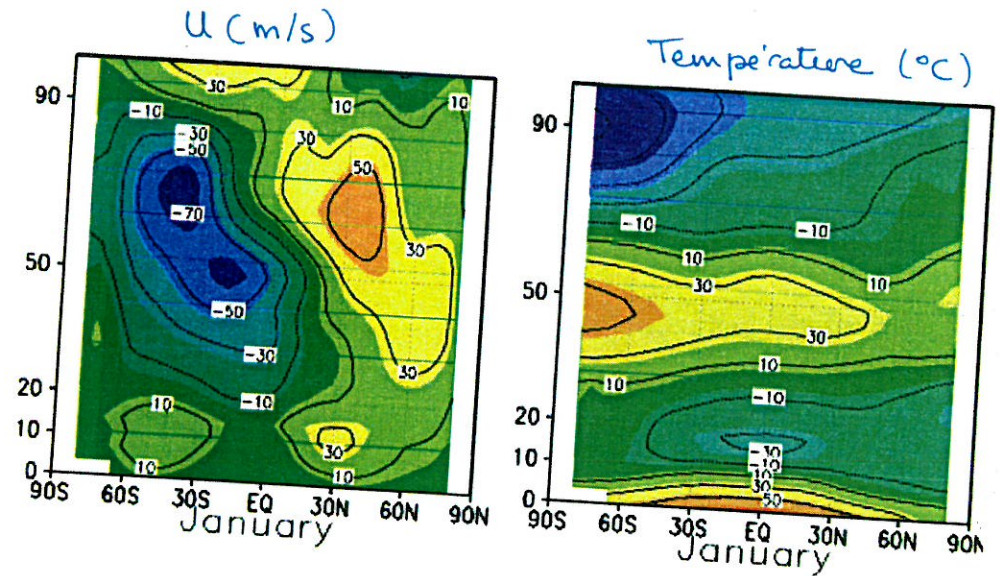
②

③



## In the mesosphere

- At 50 km in the stratosphere, there is a maximum of Temperature at the summer pole
- Thermal wind: In January,  $u > 0$  in the Northern Hemisphere  $u < 0$  in the Southern Hemisphere
- But not in the mesosphere, because of gravity wave breaking!



## 4. Why are planetary waves important for stratospheric circulation?

-At 50 km in the stratosphere, there is a maximum of Temperature at the summer pole

- Thermal wind: In January,  $u > 0$  in the Northern Hemisphere  $u < 0$  in the Southern Hemisphere

-But the meridional gradient of temperature is less strong than what is expected by radiative considerations.

- This is because of the Brewer-Dobson circulation, linked to Rossby-wave breaking.

