

Clouds

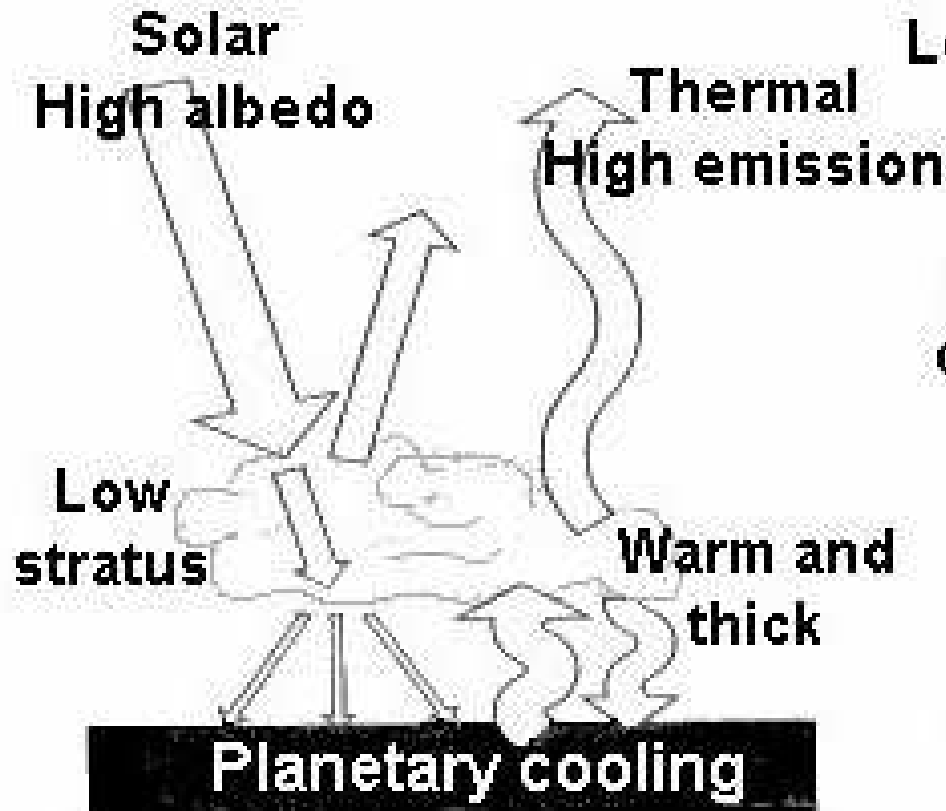
LMDz Training – December 2018
J-B Madeleine, C. Rio and the LMDz team



Picture by Oleg Artemyev taken from the ISS

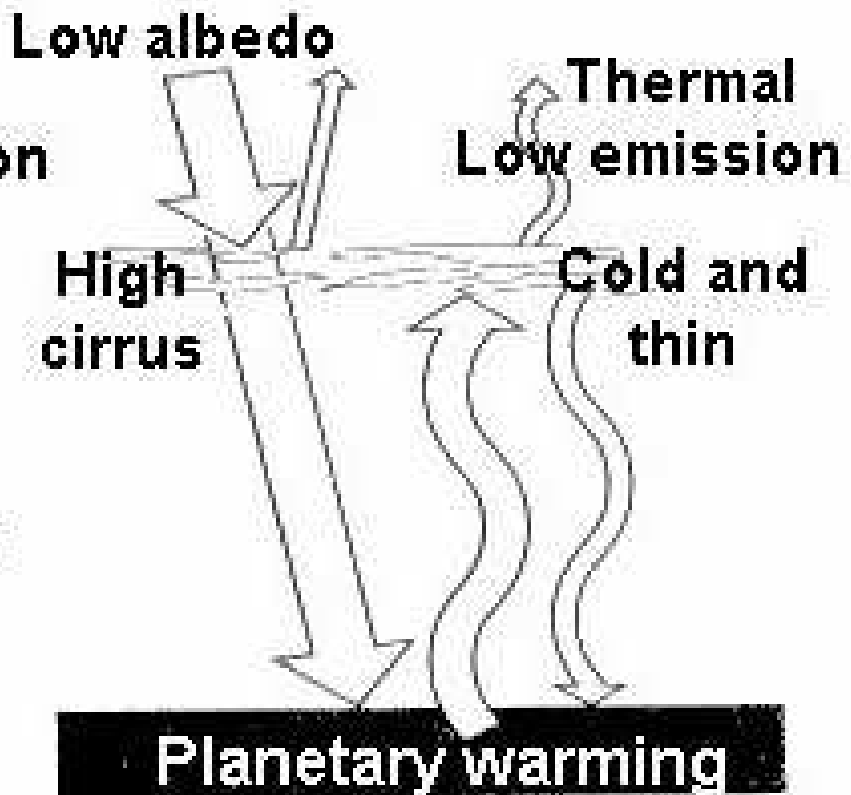
Radiative impact of clouds

(a) Low clouds



- Low clouds
- albedo effect (reflectivity of 40-50%)
 - weak greenhouse effect (high temp)

(b) High clouds



- High clouds :
- weak albedo effect
 - strong greenhouse effect (cold clouds)

Radiative forcing

LW radiative forcing

Positive : clouds reduce the LW outgoing radiation

Annual mean : $+29 \text{ W m}^{-2}$

SW radiative forcing

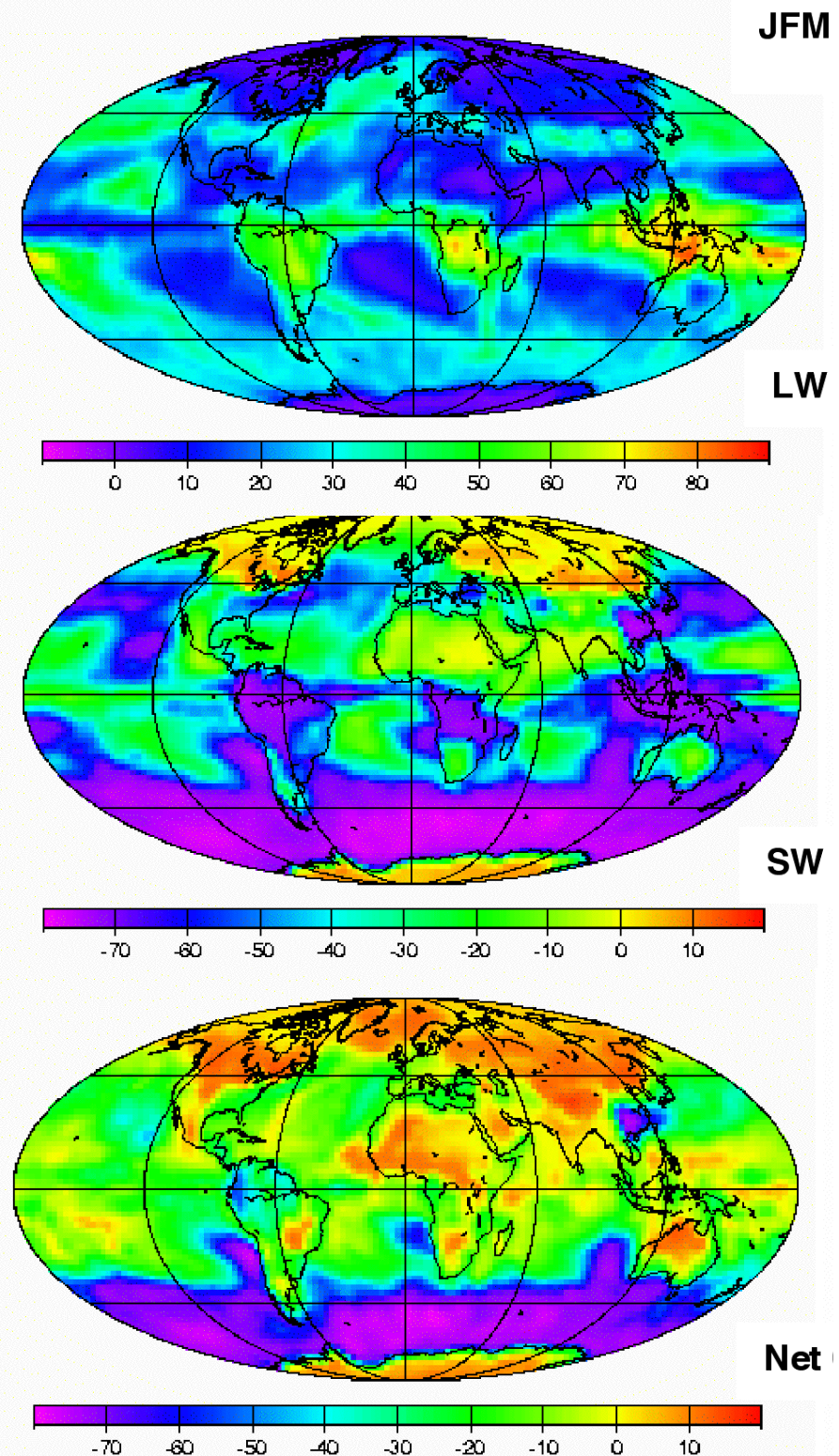
Negative : clouds reflect the incoming SW radiation

Annual mean : -47 W m^{-2}

Net forcing : **Cooling**

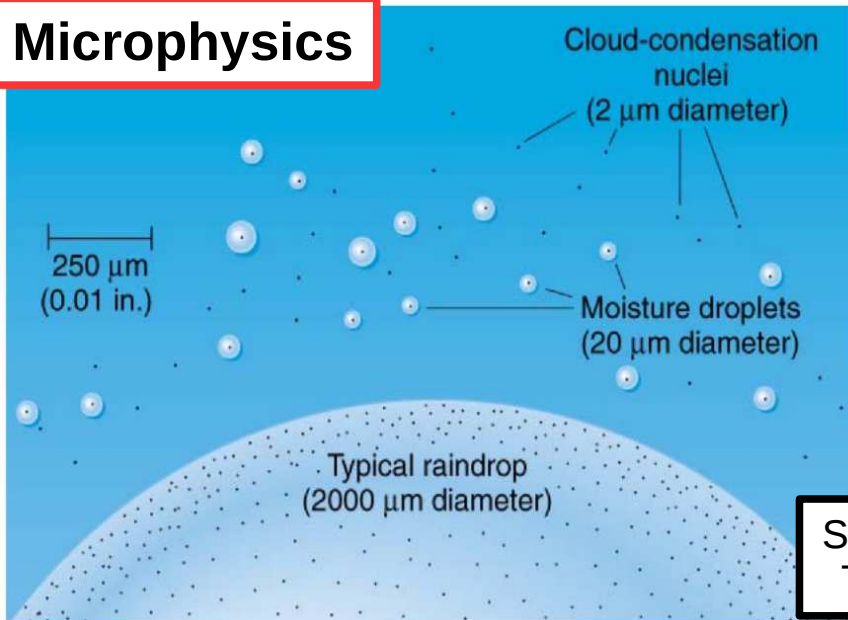
Annual mean : -18 W m^{-2}

« The single largest uncertainty in determining the climate sensitivity to either natural or anthropogenic changes are clouds and their effects on radiation » IPCC report



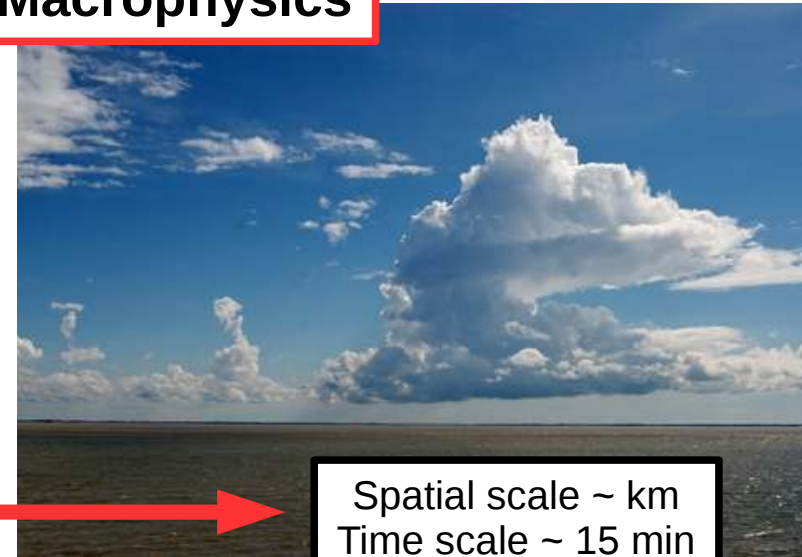
Modeling clouds : a challenge

Microphysics

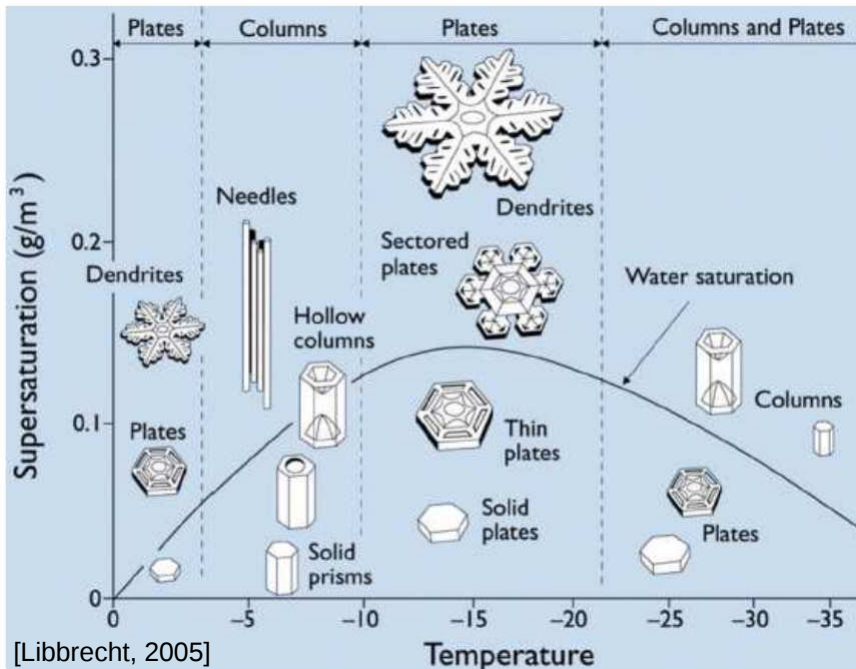


Spatial scale $\sim \mu\text{m}$
Time scale $\sim 1\text{ s}$

Macrophysics



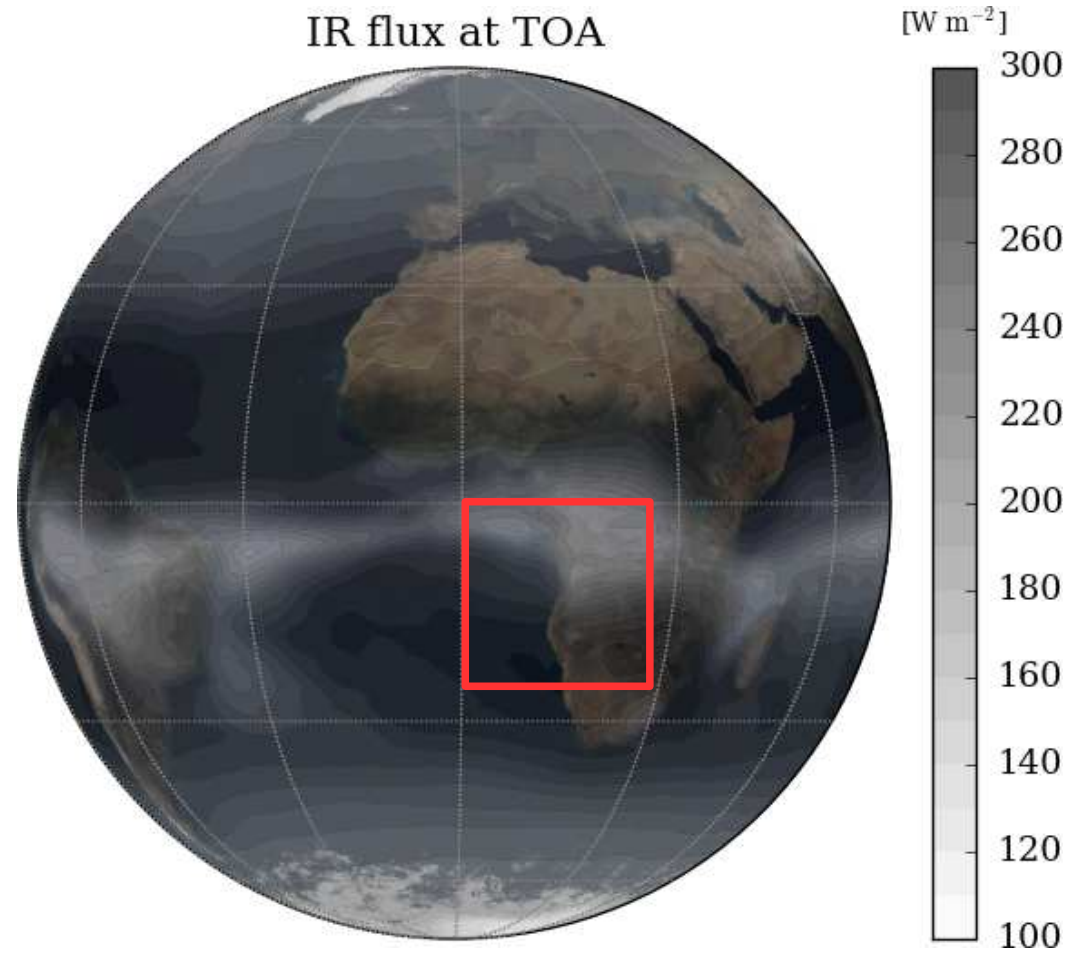
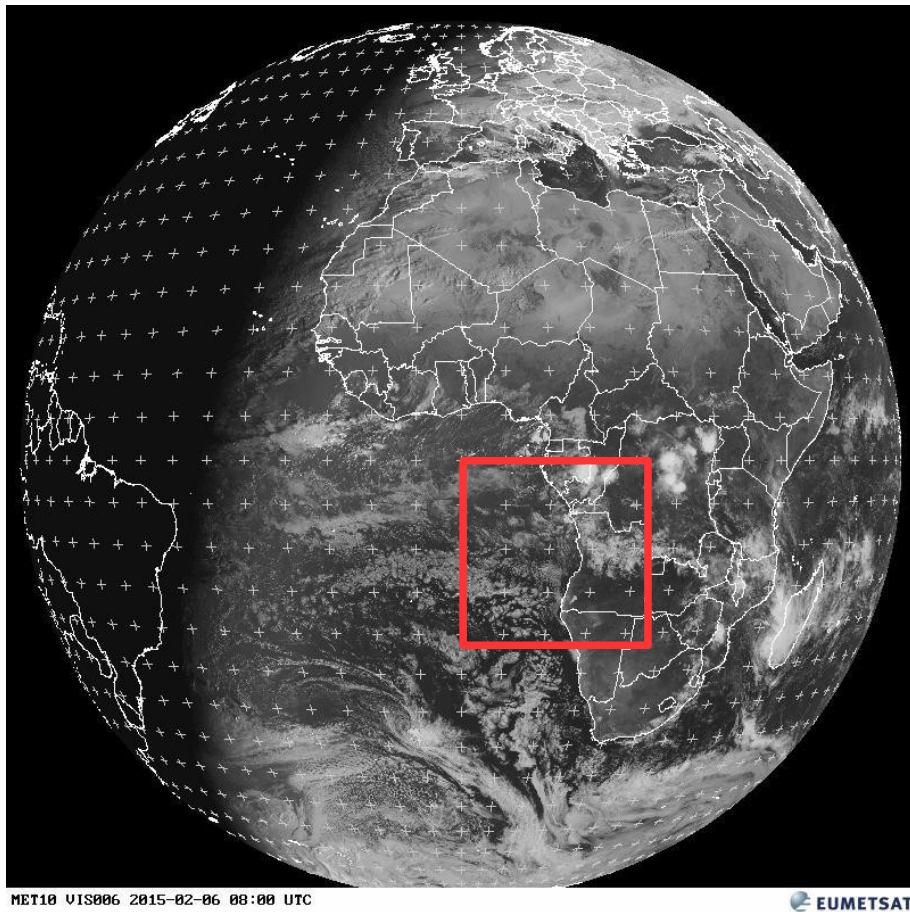
Spatial scale $\sim \text{km}$
Time scale $\sim 15\text{ min}$



[Libbrecht, 2005]



A wide variety of processes



[IPSL Climate Model / Graphisme: Planetoplot]

Fundamental process

- Clausius-Clapeyron equation :

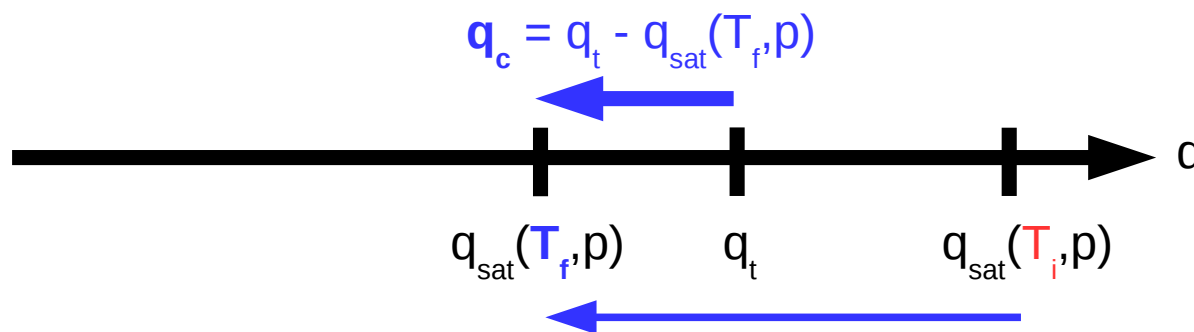
$$\frac{1}{e_{\text{sat}}} \frac{de_{\text{sat}}}{dT} = \frac{L}{R_{\text{vap}} T^2}$$

T	0°C	20°C
e_{sat}	6.1 hPa	23.4 hPa
q_{sat}	3.7 g kg ⁻¹	14.4 g kg ⁻¹

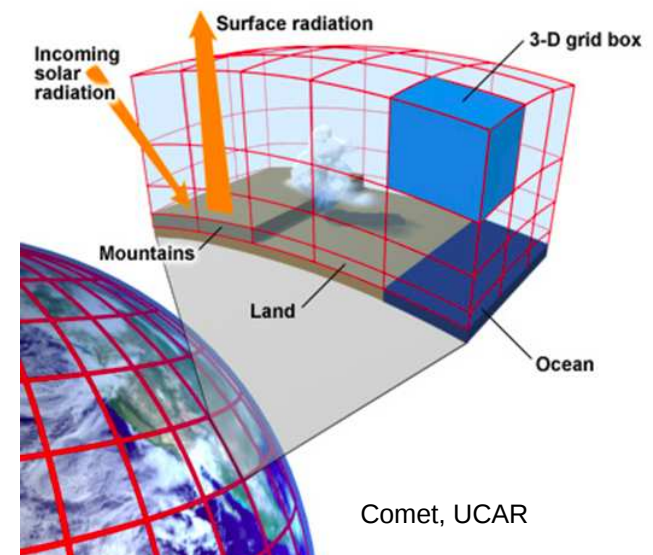
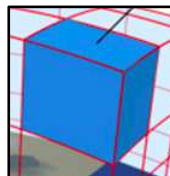
- Saturation mass mixing ratio :

$$q_{\text{sat}}(T, p) \simeq 0.622 \frac{e_{\text{sat}}(T)}{p}, \text{ where } e_{\text{sat}}(T) \text{ grows exponentially with temperature}$$

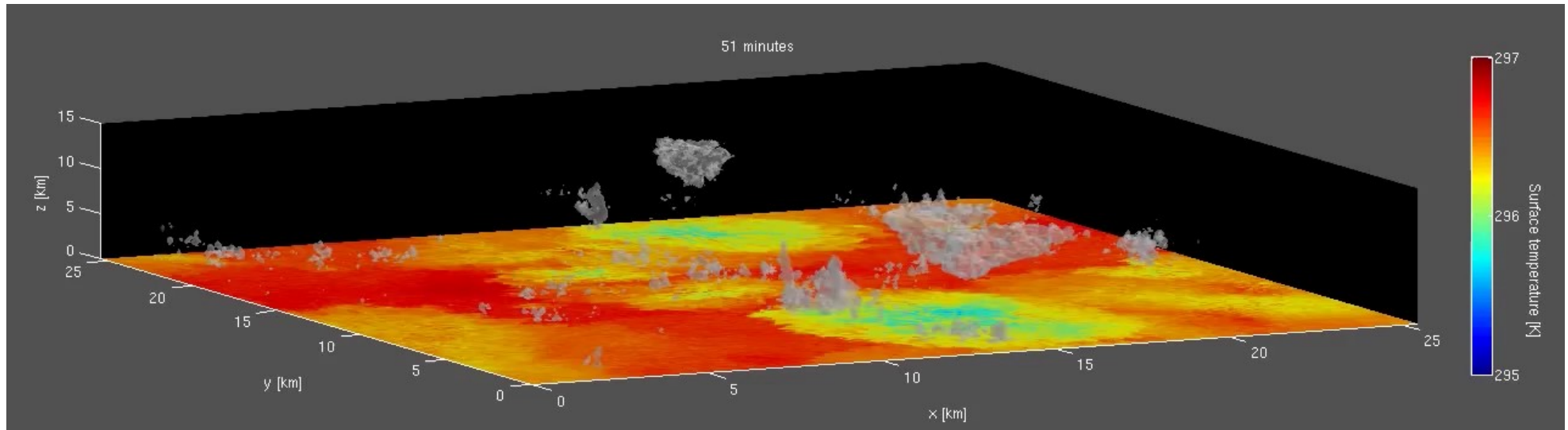
- Clouds form when an air parcel is cooled :



- But clouds do not look like that :

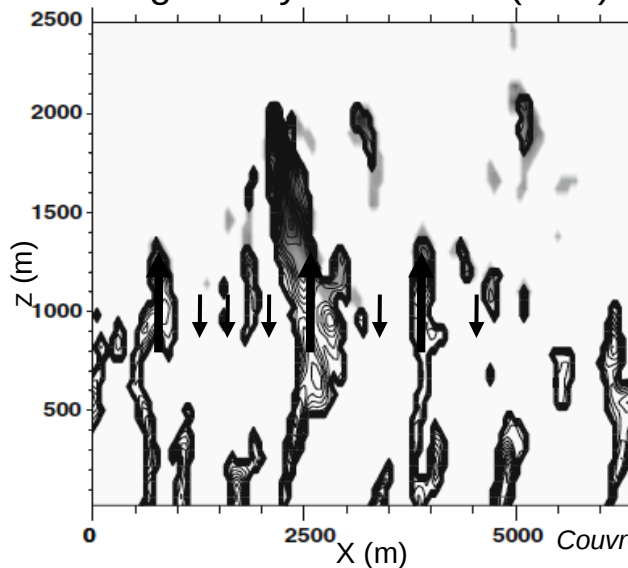


Many processes in one grid cell



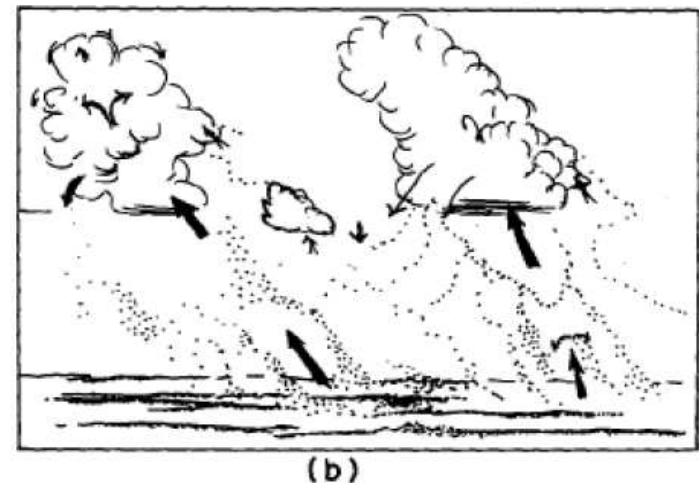
Around 8 hours of simulation by a **Cloud Resolving Model (CRM)** – C. Muller, LMD

Thermals in a Large-Eddy Simulation (LES)



Conditional sampling of thermals based on a tracer emitted at the surface.

Couvreur et al., BLM, 2010

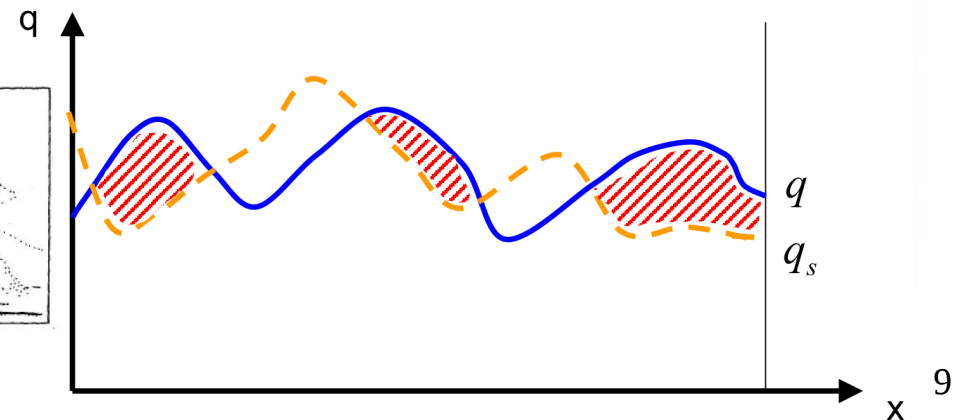
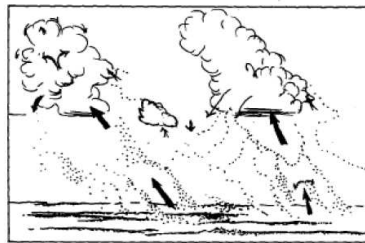
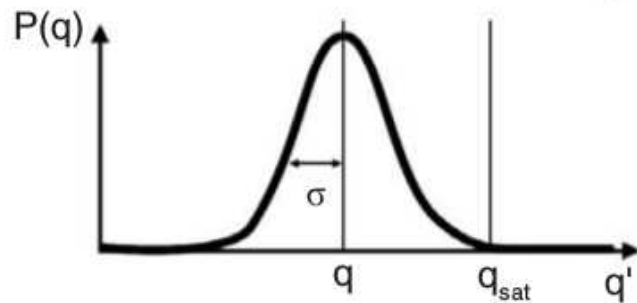
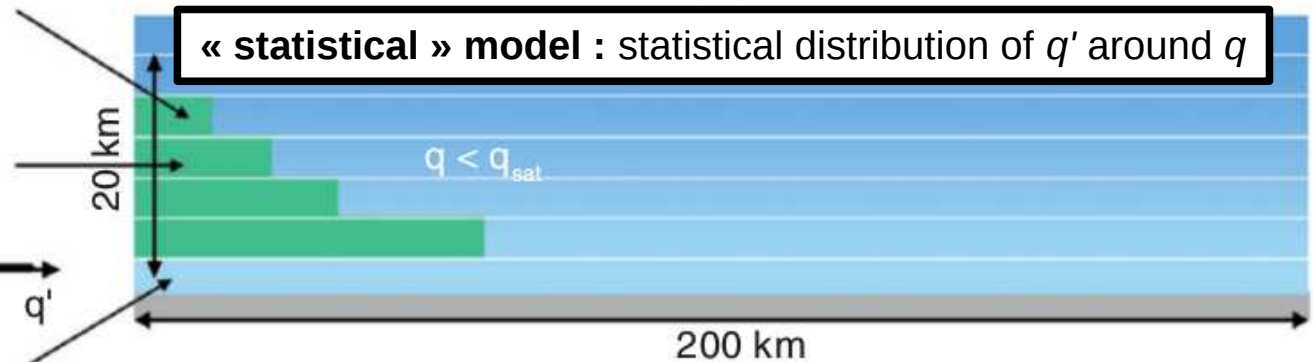
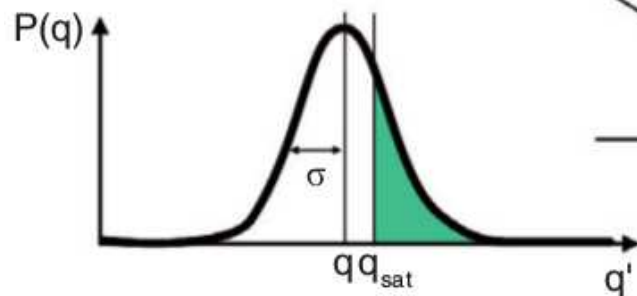
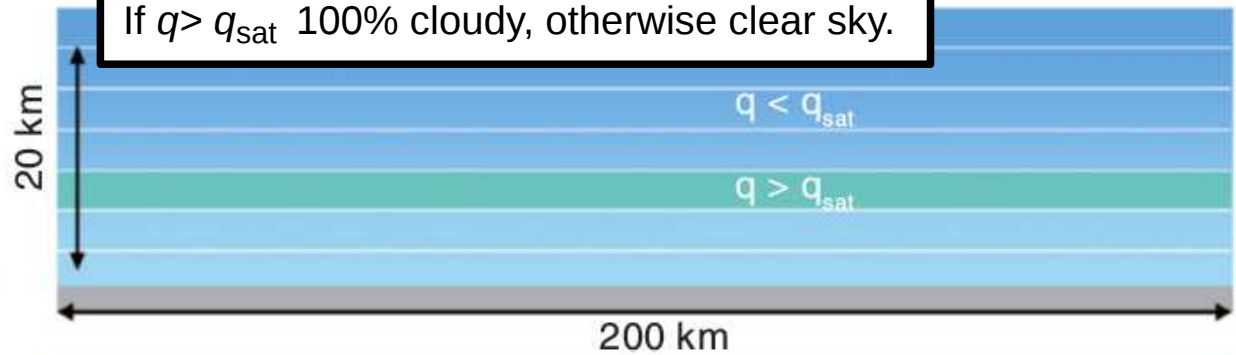
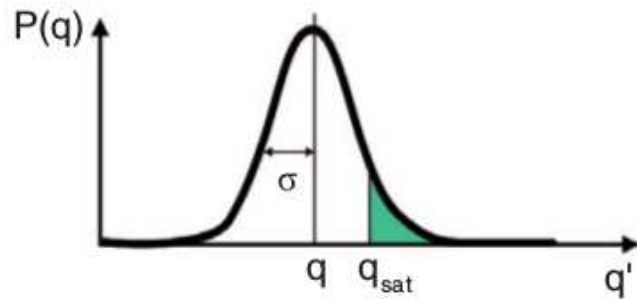


Lemone et Pennell, MWR, 1976

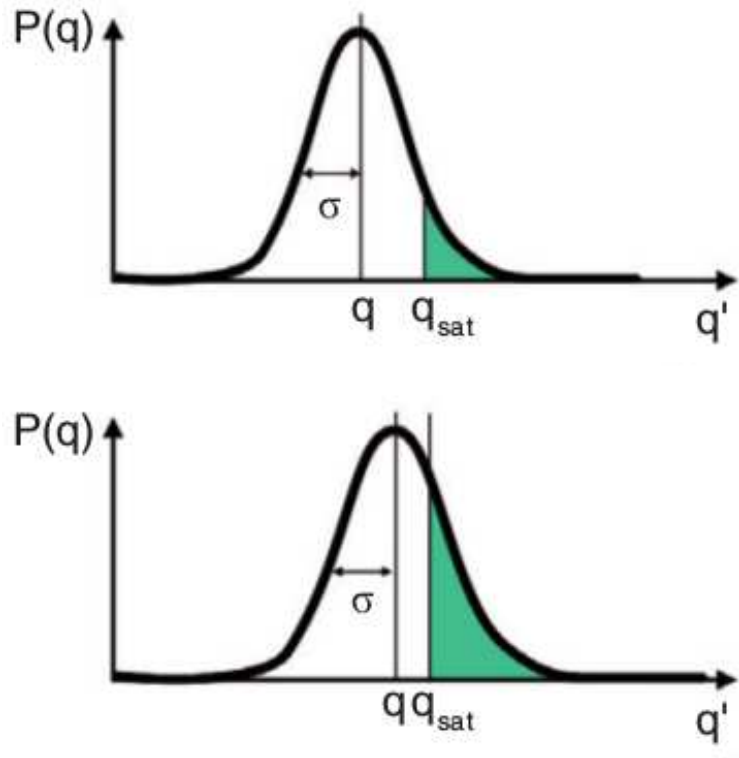
Statistical cloud scheme

« all or nothing » model :

If $q > q_{\text{sat}}$ 100% cloudy, otherwise clear sky.



Statistical cloud scheme 2/2



The goal of a cloud scheme is therefore to compute q_c^{in} and the cloud fraction based on the different physical parameterizations.

Mean total water content :

$$\bar{q} = \int_0^{\infty} q P(q) dq$$

Domain-averaged amount of condensate :

$$q_c = \int_{q_{sat}}^{\infty} (q - q_{sat}) P(q) dq$$

Cloud fraction :

$$\alpha_c = \int_{q_{sat}}^{\infty} P(q) dq$$

In-cloud condensed water content :

$$q_c^{in} = \frac{q_c}{\alpha_c}$$

physiq_mod.F90 structure - I

Initialization (once) : *conf_phys*, *phyetat0*,
phys_output_open

Beginning *change_srf_frac*, *solarlong*

Cloud water evap. *reevap*

Vertical diffusion (turbulent mixing) *pbl_surface*

Deep convection *conflx* (Tiedtke) or *concvl* (Emanuel)

Deep convection clouds *clouds_gno*

Density currents (wakes) *calwake*

Strato-cumulus *stratocu_if*

Thermal plumes *calltherm* and *ajsec* (sec = dry)

Thermal plume clouds *calcratqs*

Large scale condensation *fisrtilp*

Diagnostic clouds for Tiedtke *diagcld1*

Aerosols *readaerosol_optic*

Cloud optical parameters *newmicro* or *nuage*

Radiative processes *radlwsu* (bis)

In blue : subroutines and instructions modifying state
variables

CAREFUL : clouds are evaporated/sublimated at the beginning of each time step (~15 min), but vapor, droplets and crystals are prognostic variables. In other words, clouds can move but can't last for more than one timestep (meaning that for example, crystals can't grow over multiple timesteps).

LMDz physics parameterizations

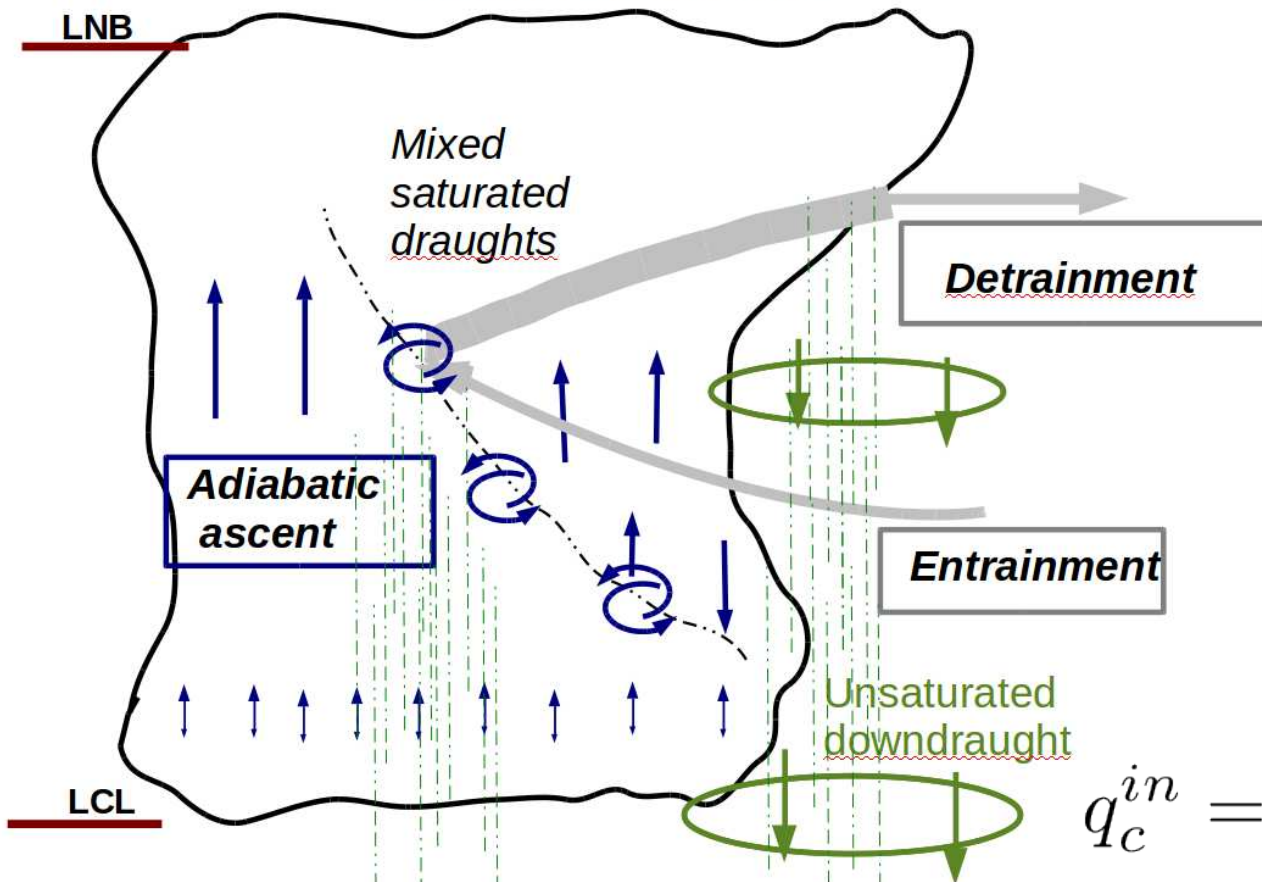
1.

2.

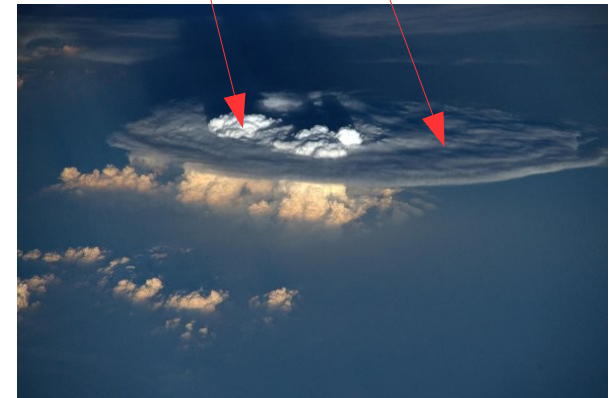
3.

1. Deep convection

Emanuel scheme



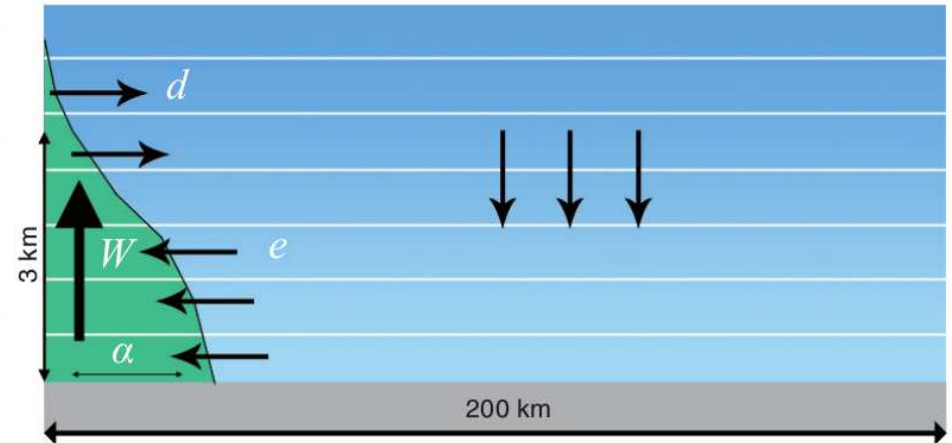
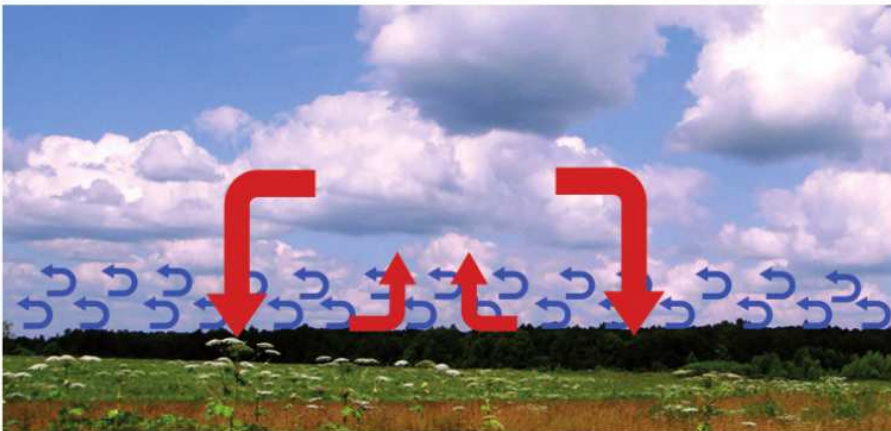
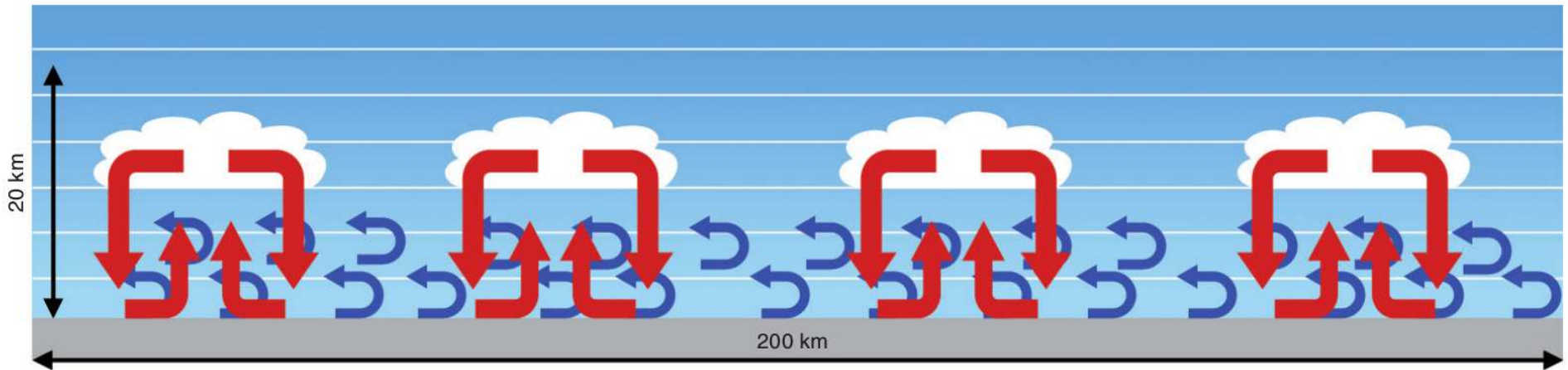
$$q_c^{in} = \frac{\sigma_a q_{ca} + \sigma_m q_{cm}}{\sigma_a + \sigma_m}$$



$$q_c^{in} = \frac{\frac{M_a}{\rho w_a} q_{ca} + \frac{\tau M_t g}{\delta p} q_{cm}}{\frac{M_a}{\rho w_a} + \frac{\tau M_t g}{\delta p}}$$

q_c^{in} is computed by the deep convection scheme and \bar{q} is known \rightarrow cloud fraction is found

2. Shallow convection 1/2



2. Shallow convection 2/2

Bi-Gaussian distribution
of saturation deficit s :

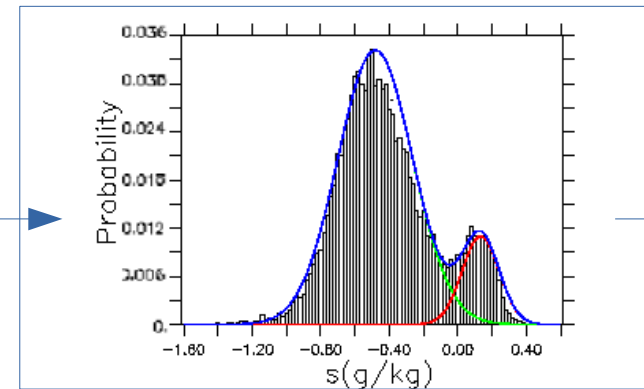
$$S = a_i (q_t - q_{\text{sat}}(T))$$

- One mode associated with thermals
 $s_{\text{th}}, \sigma_{\text{th}}$

- One mode associated with their environment:

$s_{\text{env}}, \sigma_{\text{env}}$

$s_{\text{env}}, \sigma_{\text{env}}$
 $s_{\text{th}}, \sigma_{\text{th}}, \alpha$



Jam & al., BLM, 2013

We know:

Mean state: s_{env}

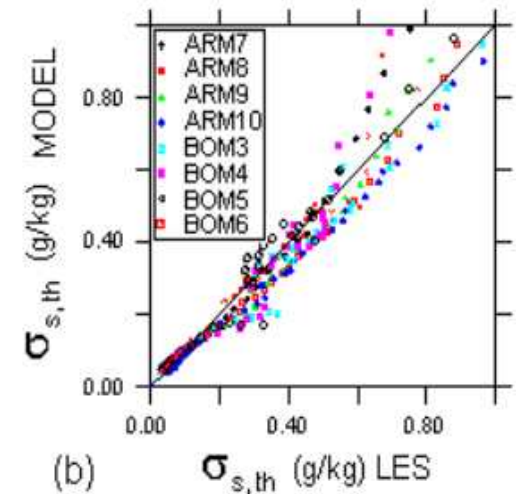
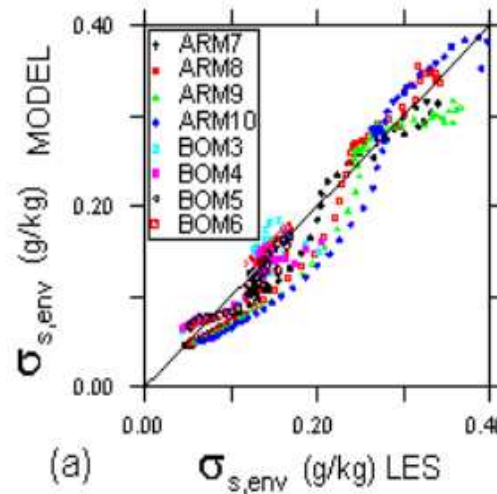
Thermal properties: s_{th}, α

Parameterization of the variances:

$$\sigma_{s,\text{env}} = c_{\text{env}} \frac{\alpha^{\frac{1}{2}}}{1 - \alpha} (\bar{s}_{\text{th}} - \bar{s}_{\text{env}}) + b \bar{q}_{t_{\text{env}}}$$

$$\sigma_{s,\text{th}} = c_{\text{th}} \alpha^{-\frac{1}{2}} (\bar{s}_{\text{th}} - \bar{s}_{\text{env}}) + b \bar{q}_{t_{\text{th}}}$$

q_c^{in} is deduced from the mean
water content of the environment
and thermals and the
parameterized spreads of the
two gaussian distributions



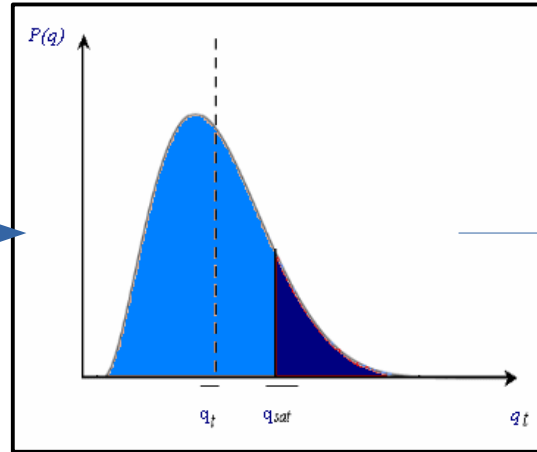
3. Large scale condensation

Log-normal distribution of total water q_t (Bony & Emanuel, JAS, 2001)

Grid-cell
mean state

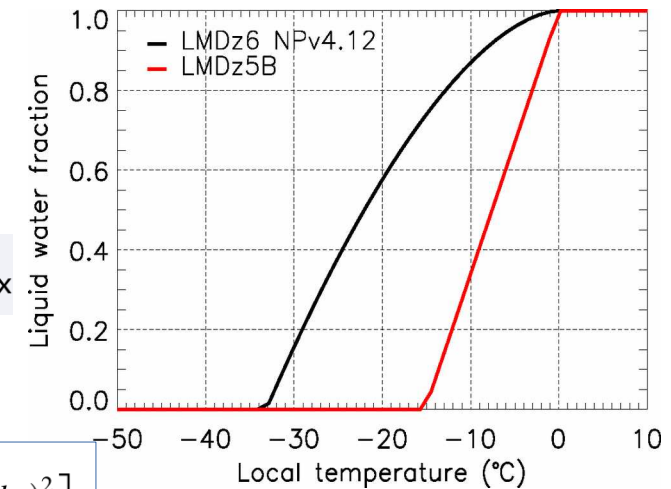
→ q, q_{sat}

σ/q imposed



$$\alpha_c = \int_{q_{sat}}^{\infty} P(q) dq$$

$$q_c = \int_{q_{sat}}^{\infty} (q - q_{sat}) P(q) dq$$



- condensate: liquid/ice partitioning (function of the temperature) :

$$^a \text{Cloud liquid fraction} = \left(\frac{T - T_{\min}}{T_{\max} - T_{\min}} \right)^n, \text{ for } T_{\min} \leq T \leq T_{\max}$$
- A fraction of the condensate falls as rain (parameters controlling the maximum water content of clouds and the auto-conversion rate) :

$$\frac{dq_{lw}}{dt} = -\frac{q_{lw}}{\tau_{convers}} \left[1 - e^{-(q_{lw}/clw)^2} \right]$$

- The rain is partly evaporated in the grid below (parameter controlling the evaporation rate) :

$$\frac{\partial P}{\partial z} = \beta [1 - q/q_{sat}] \sqrt{P}$$

$$\frac{dq_{iw}}{dt} = \frac{1}{\rho} \frac{\partial}{\partial z} (\rho w_{iw} q_{iw})$$

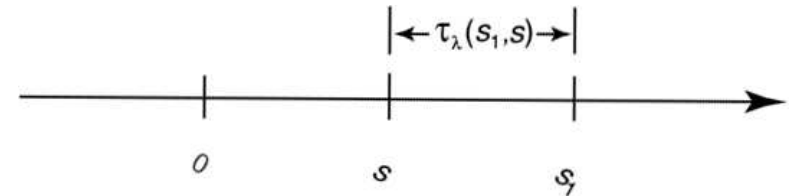
$$w_{iw} = \gamma_{iw} w_0$$

$$w_0 = 3.29 (\rho q_{iw})^{0.16}$$

Radiative transfer

Radiative transfer equation :

$$-\mu \frac{\partial I_\lambda}{\partial \tau_\lambda}(\tau_\lambda, \mu, \Phi) = -I_\lambda(\tau_\lambda, \mu, \Phi) + S_\lambda(\tau_\lambda, \mu, \Phi) + \frac{w_{0\lambda}}{4\pi} \int_0^{2\pi} \int_{-1}^1 P_\lambda(\mu, \mu', \Phi, \Phi') I_\lambda(\tau_\lambda, \mu', \Phi') d\mu' d\Phi'$$



$$\begin{aligned} q_{c,tot} = & q_c^{in} \text{ (thermals)} \times \text{CF (thermals)} \longleftarrow \text{bi-gaussian PDF} \\ & + q_c^{in} \text{ (convection)} \times \text{CF (convection)} \\ & + q_c^{in} \text{ (large-scale)} \times \text{CF (large-scale)} \longleftarrow \text{Lognormal PDF} \end{aligned}$$

$$CF_{tot} = \min(CF \text{ (thermals)} + CF \text{ (convection)} + CF \text{ (large-scale)}, 1.)$$

Solving the radiative transfer equation requires :

- $q_{c,tot}$ to compute the opacity ;
- **Cloud droplet and crystal sizes** to compute the optical properties ;
- CF_{tot} to compute the heating rates in the clear-sky ($1 - CF_{tot}$) and cloudy (CF_{tot}) columns.

Optical properties of ice crystals

(for droplets, see O. Boucher's talk)

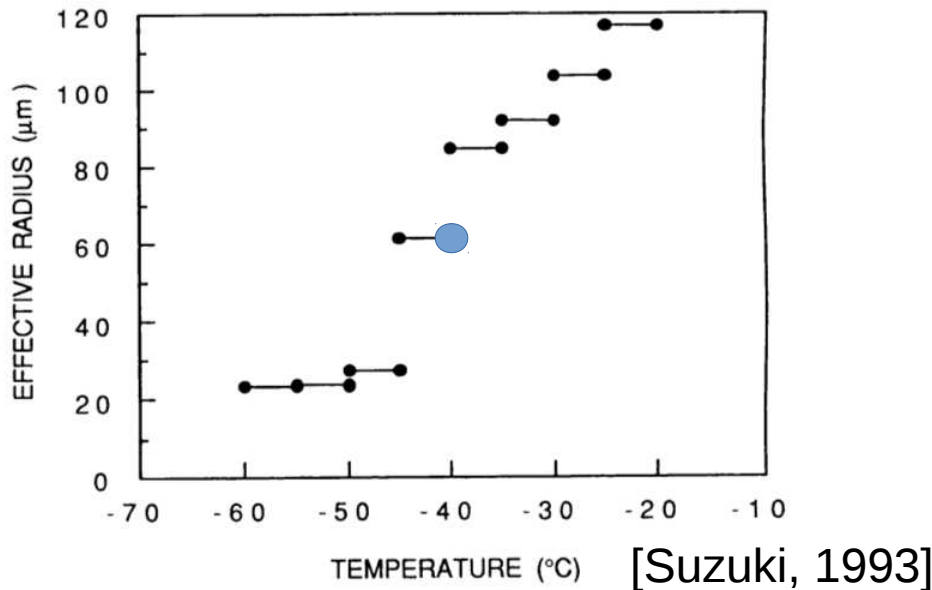


Fig. 2. Cloud temperature versus effective radius of cloud particle.

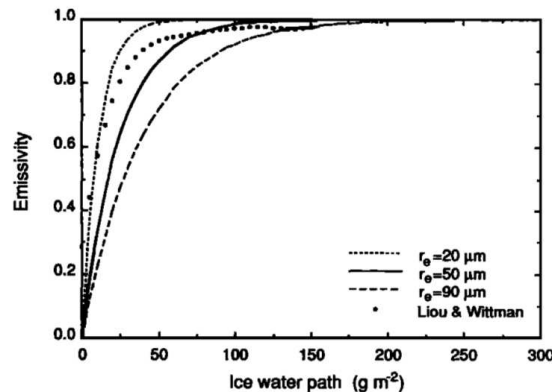
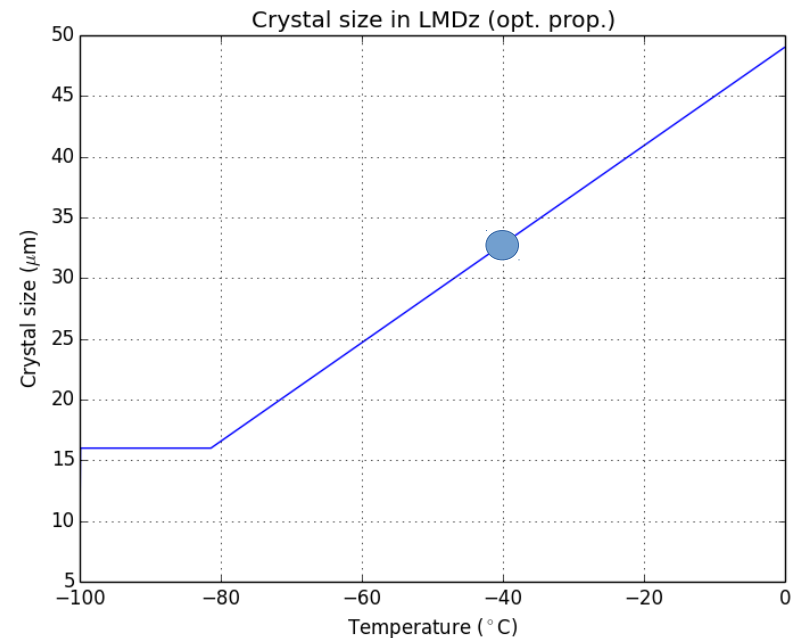


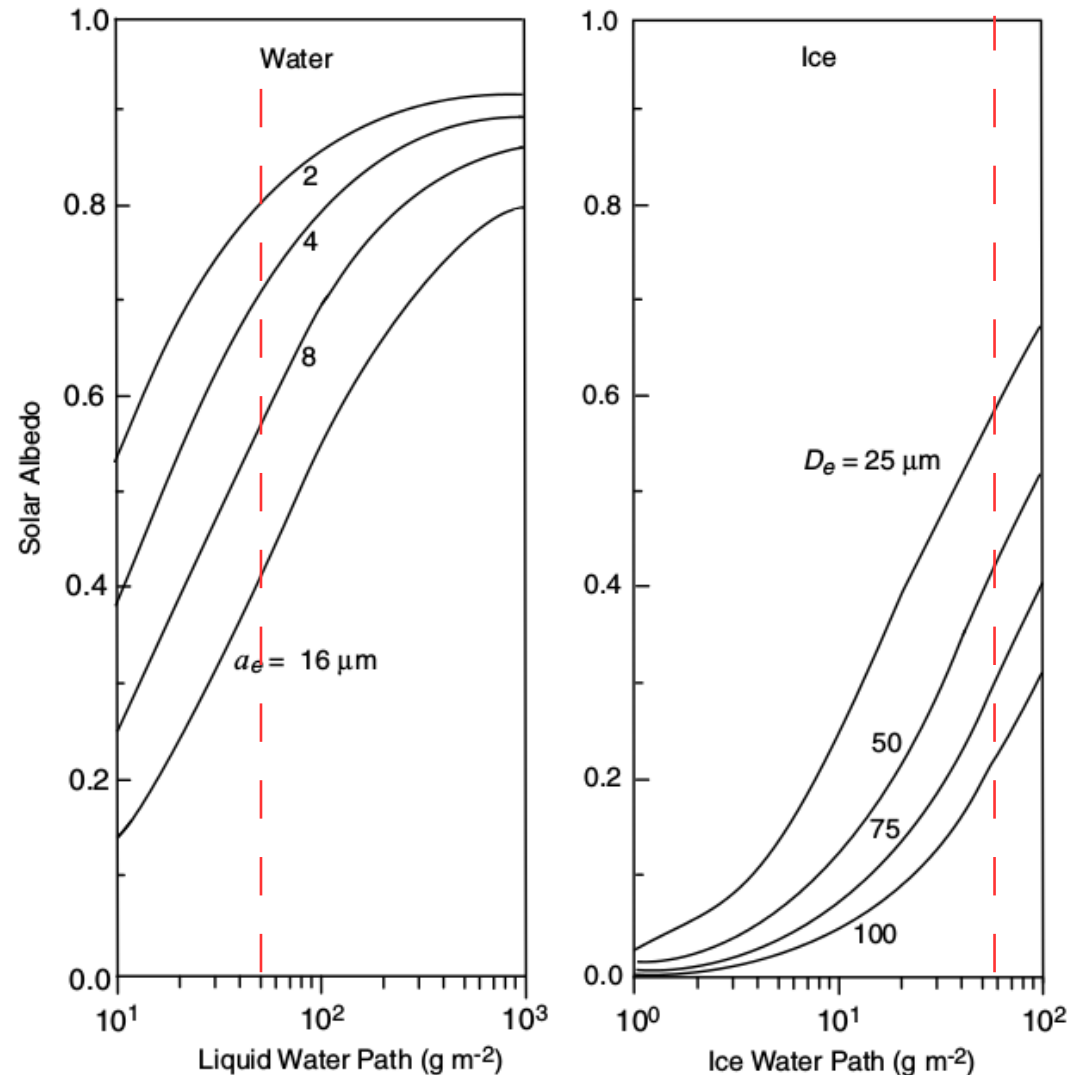
Fig. 5. Cirrus infrared emissivity for $r_e = 20, 50$, and $90 \mu\text{m}$ as a function of ice water path. The solid circles represent values computed using the parameterization of Liou and Wittman [1979].



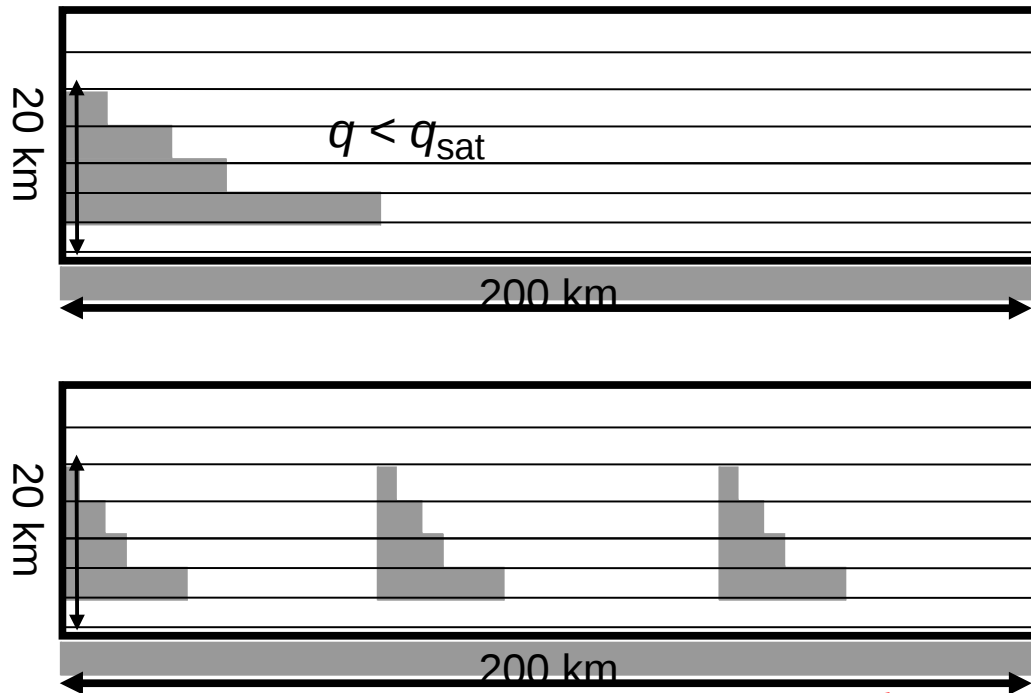
Droplet sizes follow
 $r = 0.71T + 61.29$ in μm
[Iacobellis et Somerville 2000]
 with $r_{\min} = 3.5 \mu\text{m}$ for
 $T < -81.4^\circ\text{C}$ *[Heymsfield et al. 1986]*

Importance of cloud phase

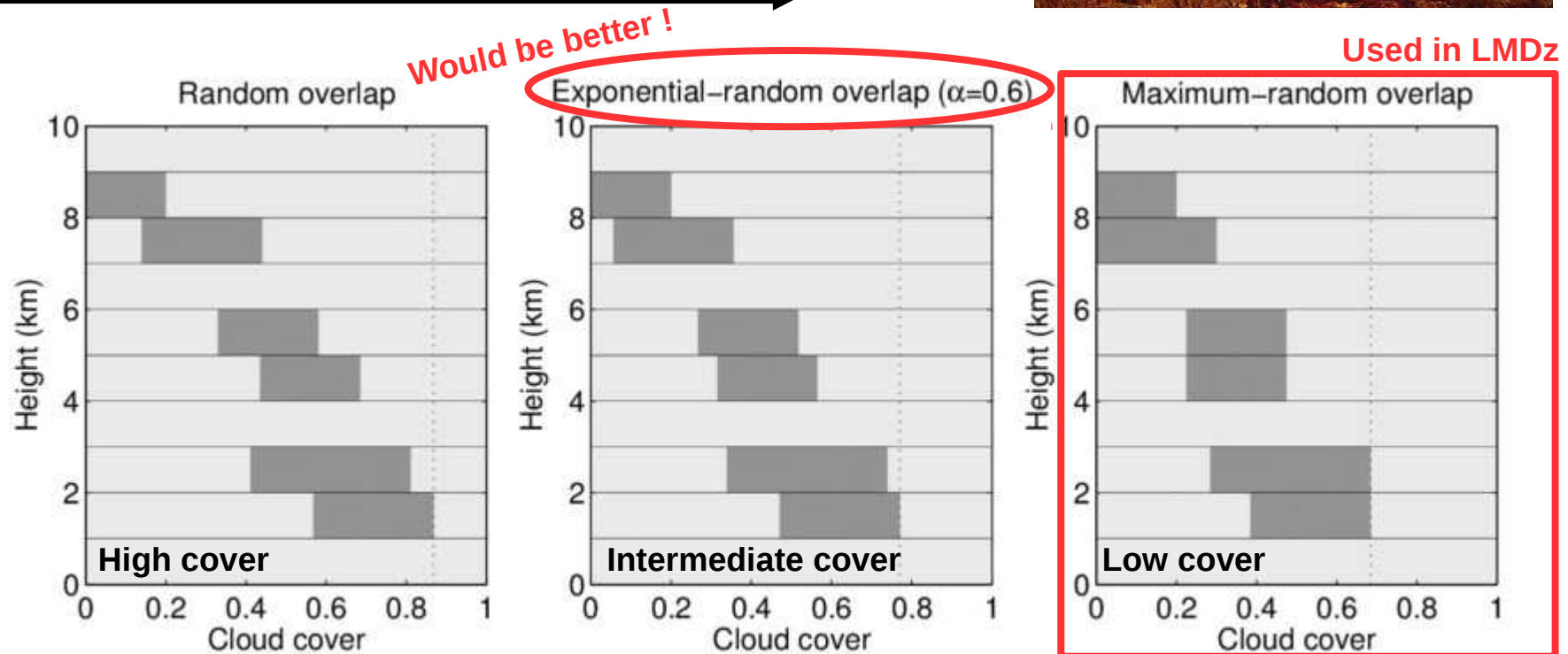
- Clouds reflect sunlight (negative forcing, cooling) and emit in the infrared (positive forcing, warming) ;
- For the same water content, liquid clouds reflect more sunlight than ice clouds ;
- For liquid clouds : if the cloud water content increases, there is a negative forcing (reflection dominates) ;
- For ice clouds : if the cloud water content increases, the forcing depends on the size of the crystals.

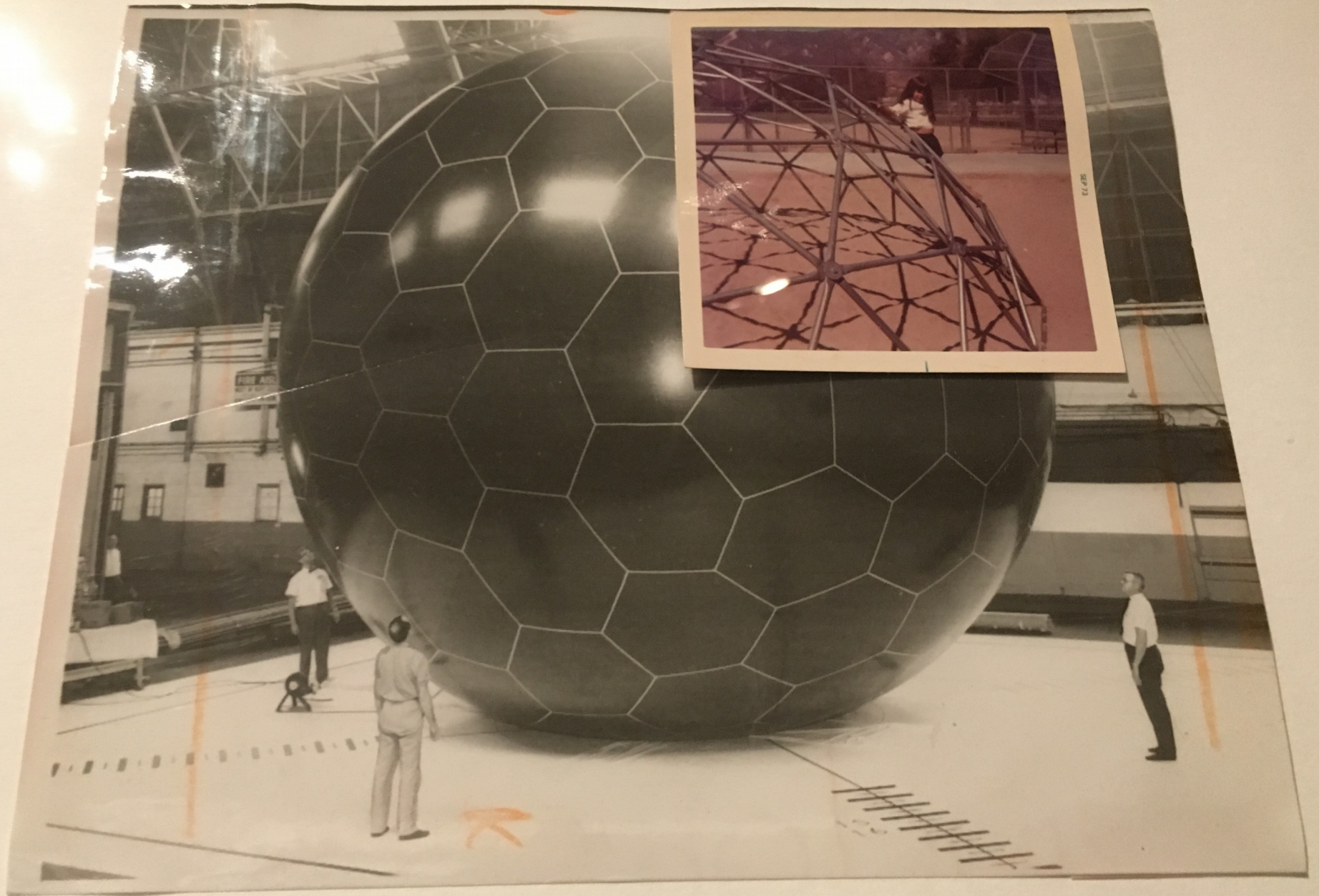


CF versus height is known, but radiation also needs to know the total cloud **cover** ; we therefore parameterize the **cloud overlap**



▲ **LMDz : Maximum random overlap**
For the GCM, these two scenes are identical ;





Welcome to the LMDz team ;-)