



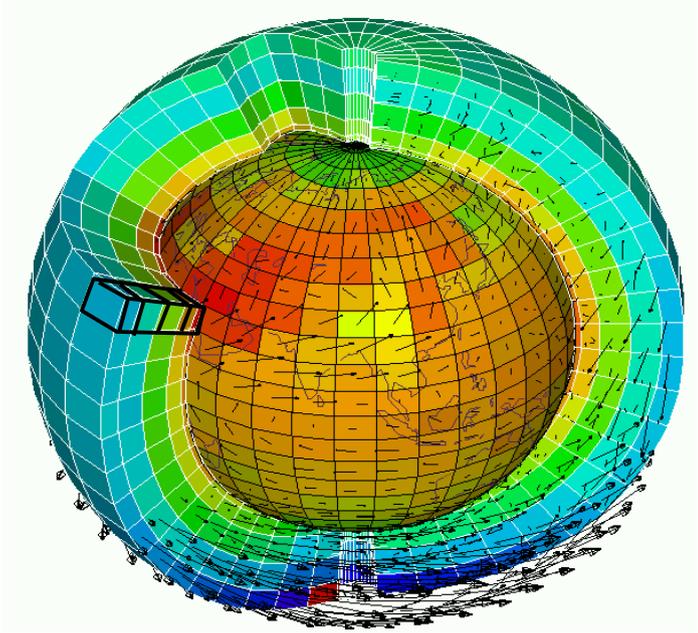
Strengths and weaknesses of the different physical packages of LMDZ

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LMDZ Development Team

LMDZ training course – 7/8/9 December 2015 – LMD Jussieu

Physical parameterizations

The general circulation model LMDZ



Dynamical component:

Discretization of conservation equations on the sphere

Physical component:

Represents the effects of sub-grid scale processes non resolved by the dynamical equations (source terms) in each atmospheric column

Key for the representation:

- vertical profiles of **heating rates** and the large-scale circulation
- Vertical profiles of **mass fluxes** and tracer transport
- Clouds and their **radiative impacts**
- **precipitation**

Key role in the **coupling** with the other components:

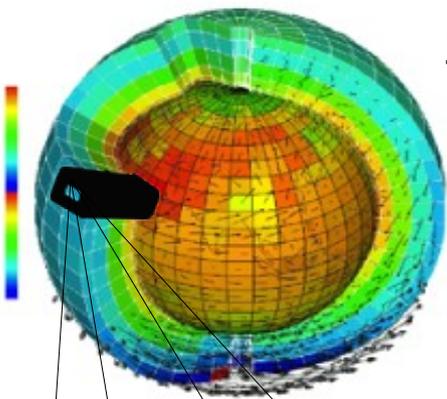
- transport of chemical species
- surface fluxes
- hydrology

Provide the variables necessary for **impact** studies:

- precipitation
- sub-grid scale variations of winds, etc...

Methodology to develop and evaluate parameterizations

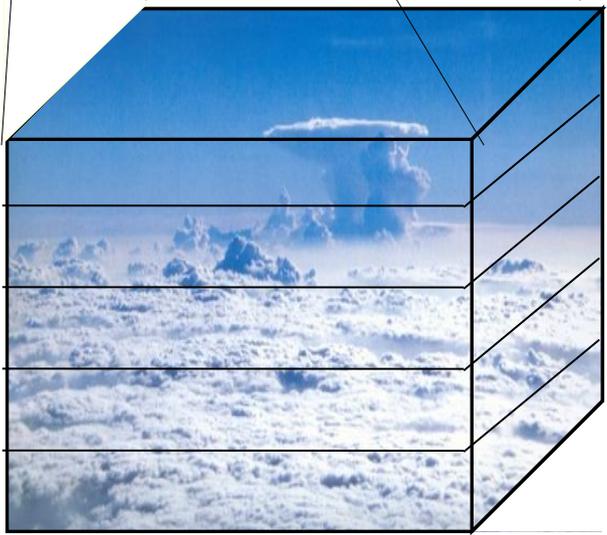
1D case studies built from field campaigns (BOMEX, TOGA-COARE, TWP-ICE, AMMA...) or routinely in-situ measurements (ARM)



1D simulations

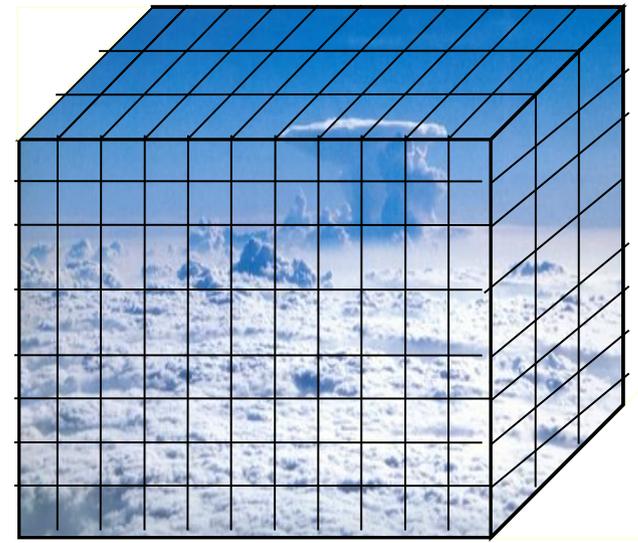
Explicit simulations over a domain equivalent to a GCM grid cell

Provide quantities difficult to measure (structures properties, mixing rates etc...)



200km

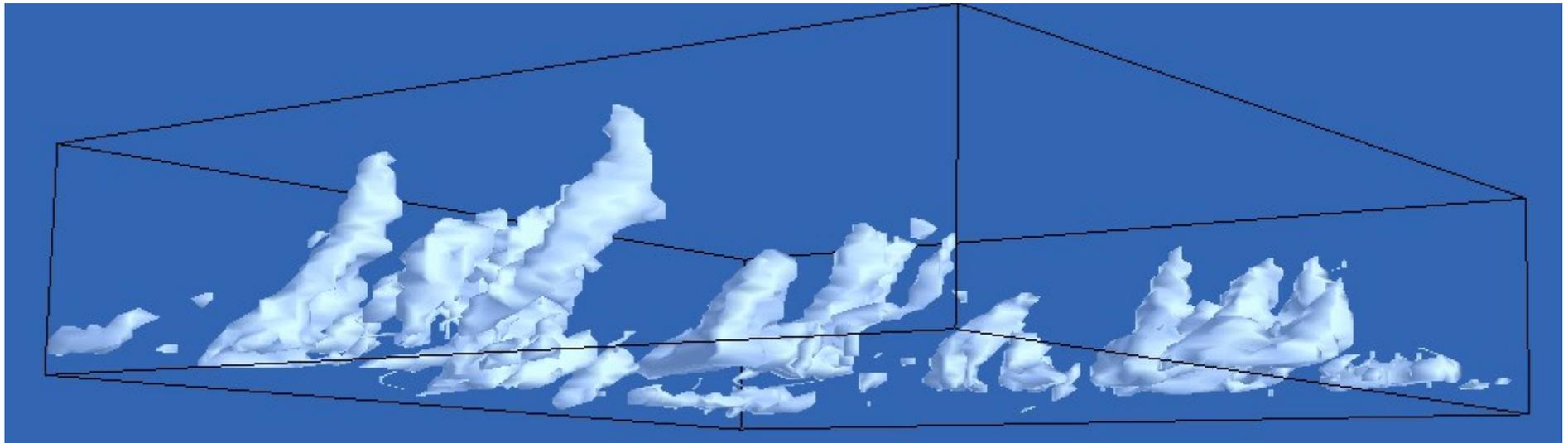
←→
Identical forcing:
Initial profiles
Large-scale advection
Surface fluxes



200km

Use of explicit simulations for parameterization development

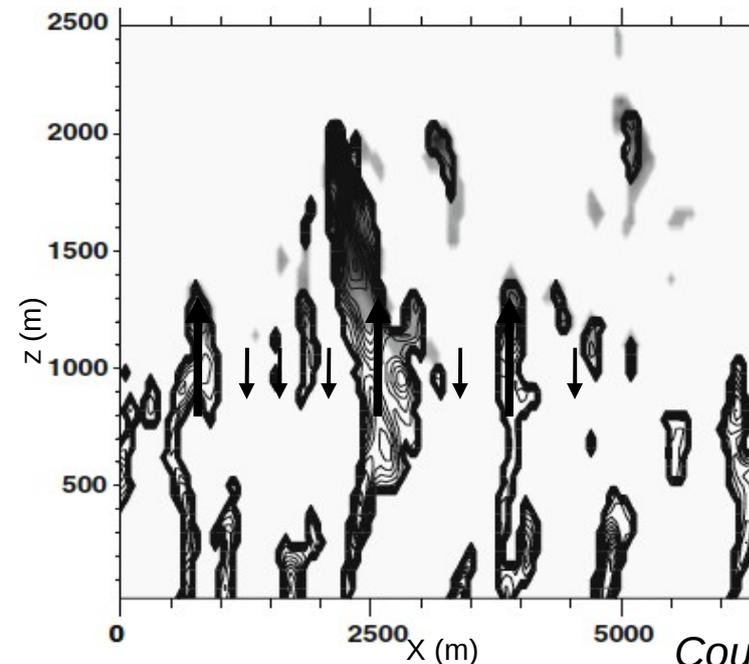
Simulated cumulus field:



<http://www.knmi.nl/~siebesma/BLCWG/>

- Evolution of mean variables:
Ex: T , q , cloud fraction (cf)
- Statistics over the domain:
Ex: PDF of qt , θ_l
- Properties of clouds:
Ex: condensate

Identification of thermals in the
Large-Eddy Simulation



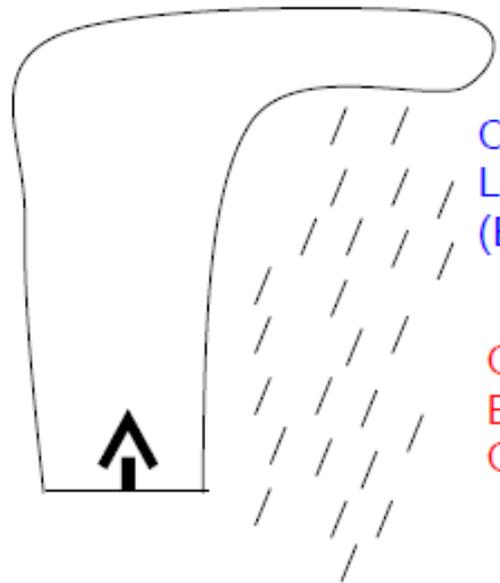
Conditional sampling
of thermals based
on a tracer emitted
at the surface.

Couvreux et al., BLM, 2010

The physical packages used for CMIP5

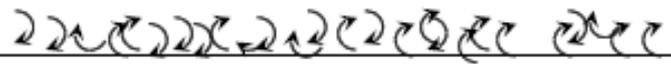
LMDZ5A

Boundary-layer:
Diffusion scheme (Louis, 1979)
(stationary equation for TKE)



Clouds:
Lognormal PDF of qt for all clouds
(Bony et Emanuel, 2001)

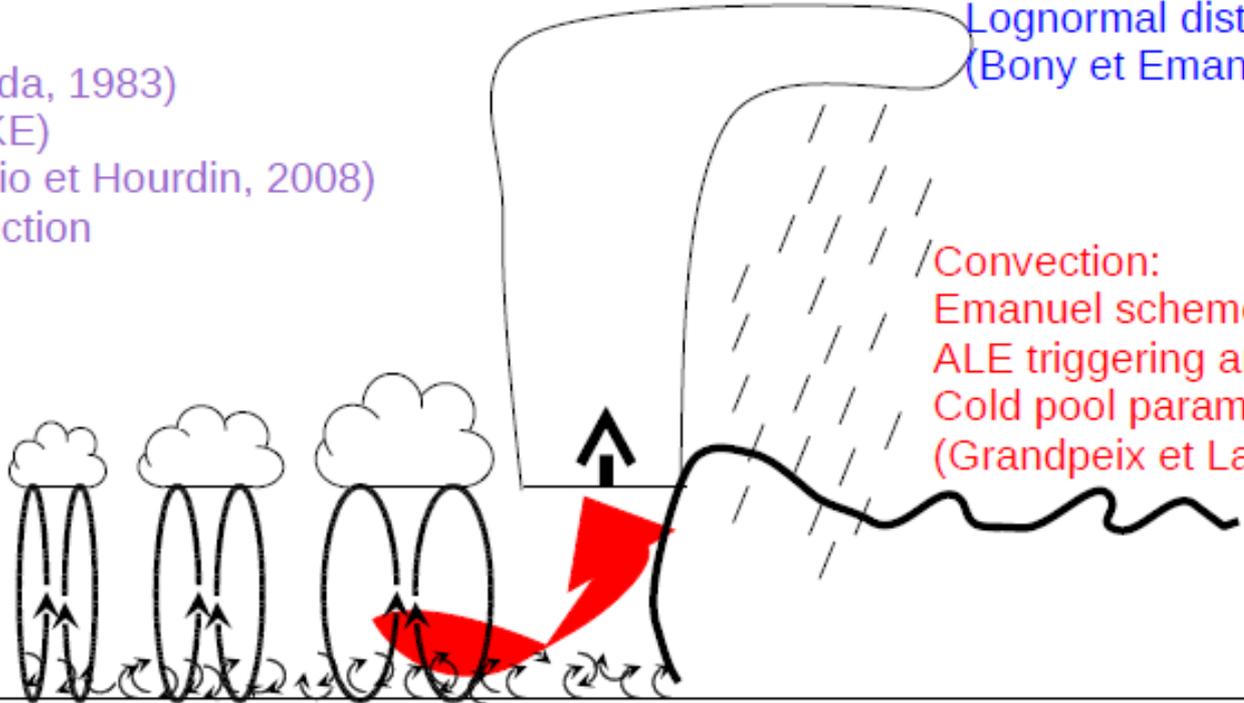
Convection:
Emanuel scheme (1991) with
CAPE closure



LMDZ5B

Boundary-layer:
- diffusion scheme (Yamada, 1983)
(pronostic equation for TKE)
- thermal plume model (Rio et Hourdin, 2008)
for dry and shallow convection

Shallow clouds:
bi-gaussian distribution
of saturation deficit
(Jam et al., 2012)



Convective and large-scale clouds:
Lognormal distribution of qt
(Bony et Emanuel, 2001)

Convection:
Emanuel scheme (1991) with
ALE triggering and ALP closure
Cold pool parameterization
(Grandpeix et Lafore, 2010)

A satellite view of Earth's atmosphere showing a dense layer of clouds over a dark, textured surface, likely the ocean. The clouds are organized into a grid-like pattern, indicating convection cells. The Earth's curvature is visible at the top of the frame.

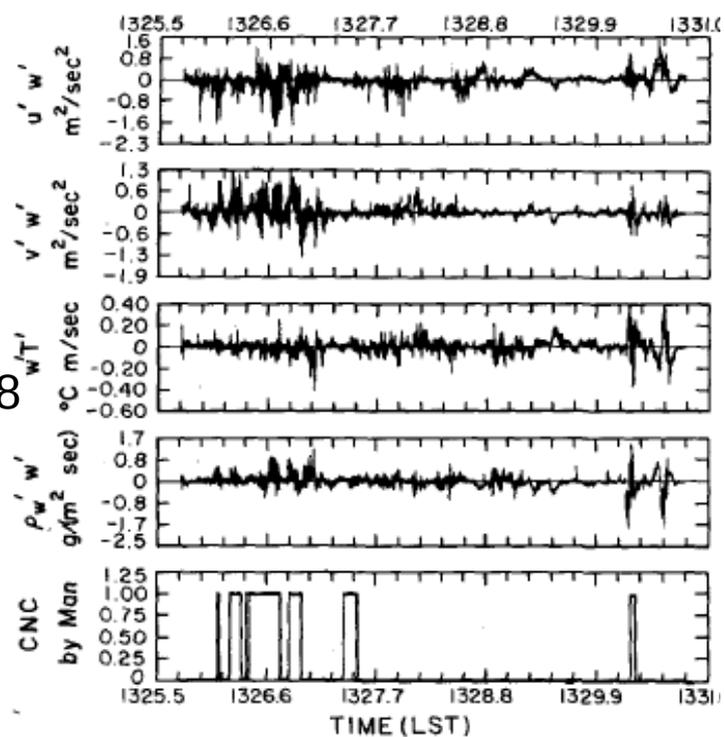
Boundary-layer convection and clouds

Cumulus and thermals

Case II: 15 dec 1972- 12h48



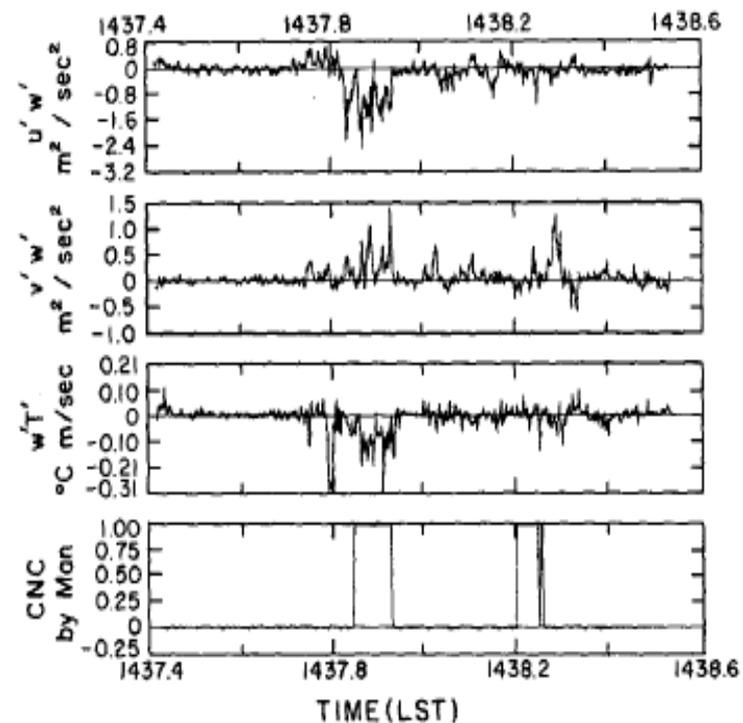
CASE II
SUPPRESSED AREA - 15 DECEMBER 1972
610 m CROSSWIND



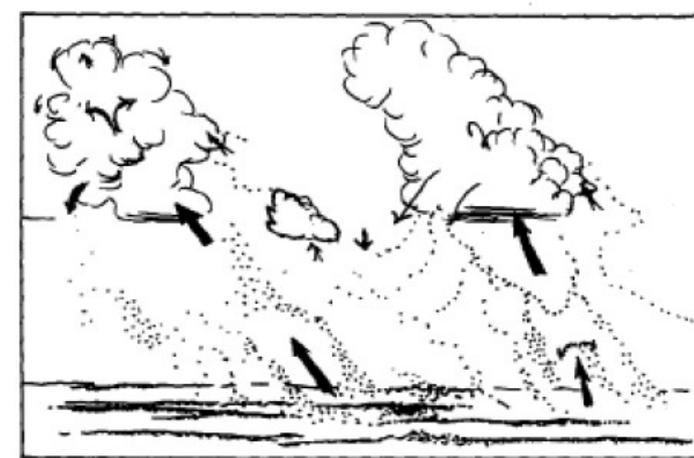
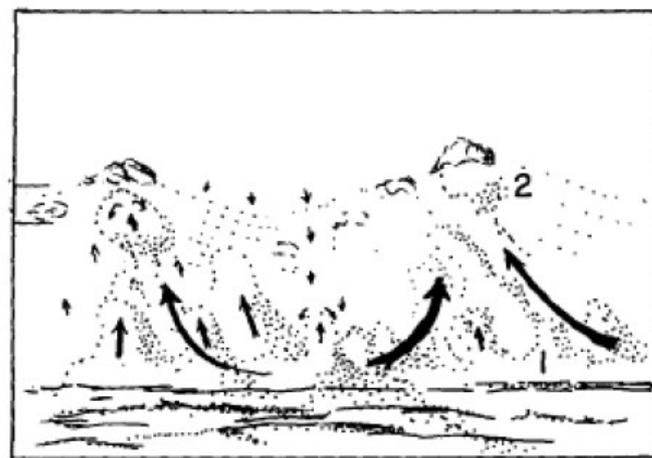
Case III: 15 dec 1972 - 14h18



CASE III
TRADE WIND CUMULUS AREA - 15 DECEMBER 1972
531 m CROSSWIND



Cumulus are the saturated part of thermals initiated at the surface

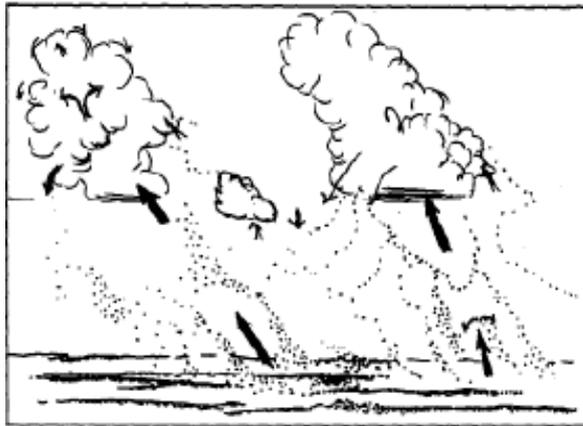


The thermal plume model

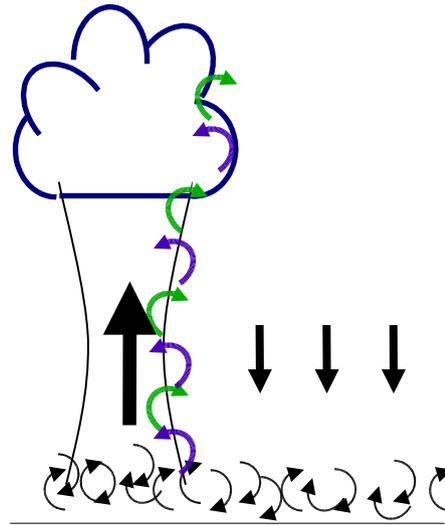
Hourdin et al., JAS, 2002; Rio et Hourdin, JAS, 2008

calltherm.F90

LeMone and Pennell, MWR, 1976



z (m)



α

$1-\alpha$

Equations

Conservation of mass:

$$\frac{\partial f}{\partial z} = e - d$$

Transport of θ_l , q_t , u , v

$$\frac{\partial f \psi_u}{\partial z} = e \psi - d \psi_u$$

Internal variables

- w : mean vertical velocity within thermals
- α : fractional coverage of thermals
- e : entrainment rate within thermals
- d : detrainment rate from thermals
- q_a : concentration of q within thermals

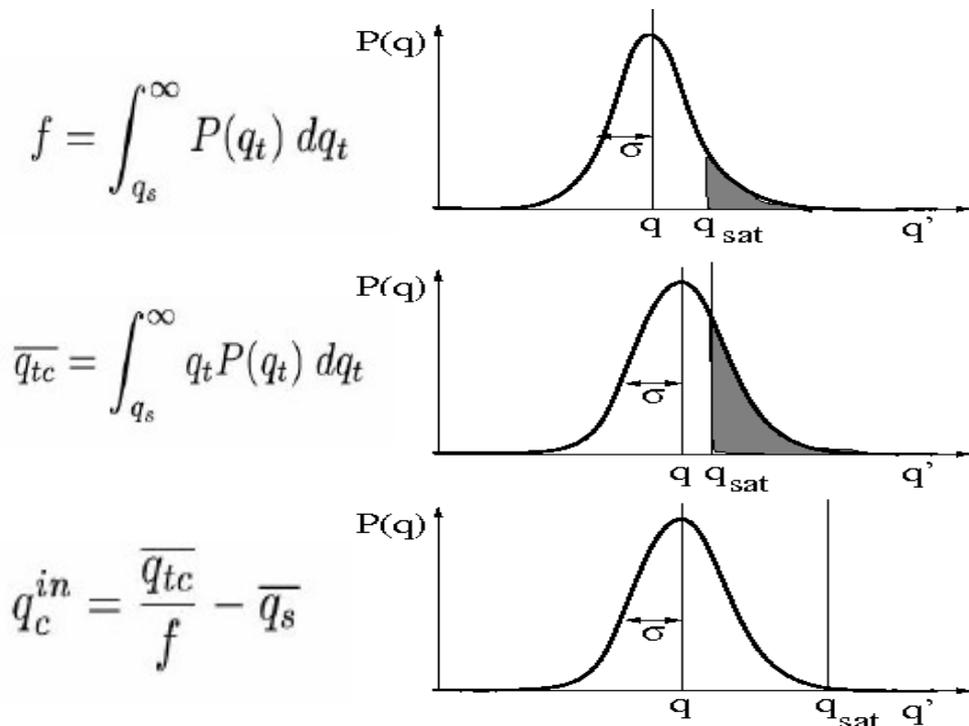
Conservation of momentum:

$$\frac{\partial f w_u}{\partial z} = -d w_u + \alpha g \rho \frac{\theta_{vu} - \theta_v}{\theta_v}$$

- + Specification of entrainment and detrainment rates
- + Computation of the mass-flux at the base of plumes

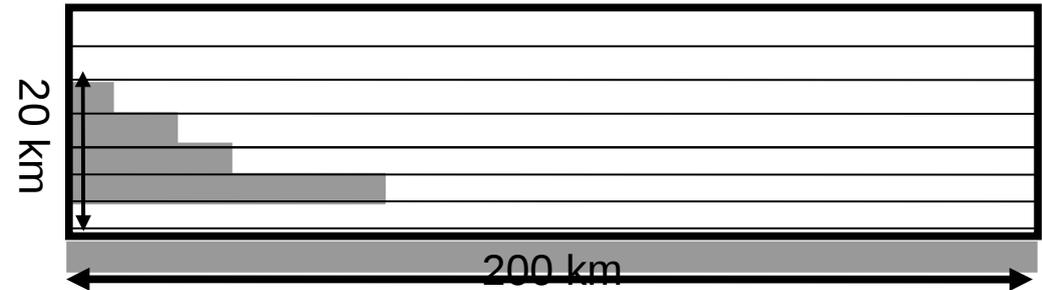
The boundary-layer cloud scheme

cloudth.F90



« statistical » model :

We assume a statistical distribution of q' around q within the grid cell



Simple parameterization : Gaussian $\sigma / q = 20\%$

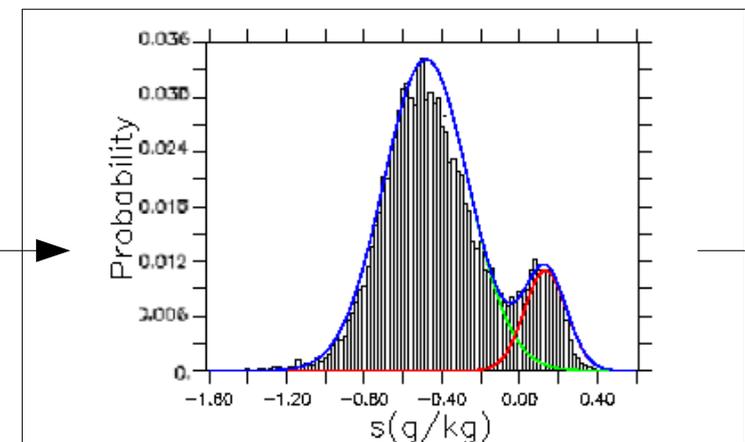
Bi-Gaussian distribution of saturation deficit s :

$$s = \alpha l (q_t - q_{sat}(T_l))$$

- One mode associated with thermals s_{th}, σ_{th}
- One mode associated with their environment: s_{env}, σ_{env}

s_{env}, σ_{env}
 $s_{th}, \sigma_{th}, \alpha$

Shallow convection



q_c, c_f

Representation of low clouds in LMDZ5A and LMDZ5B

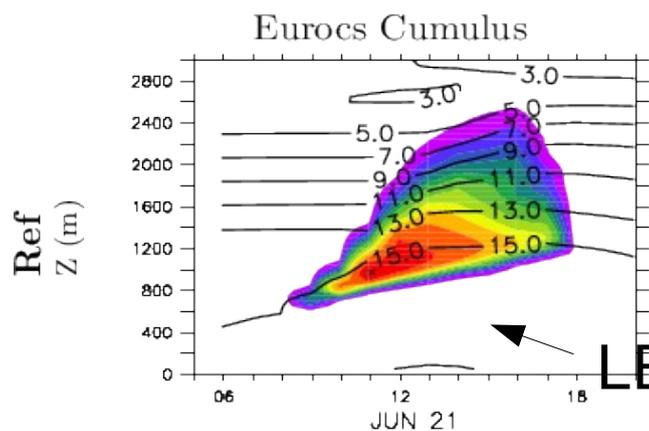
1D cases

Cloud fraction (%) and liquid water (g/kg)

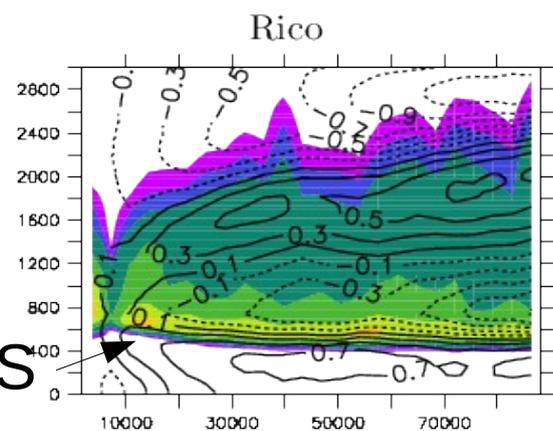
3D simulations

Low cloud fraction (%)
Annual mean

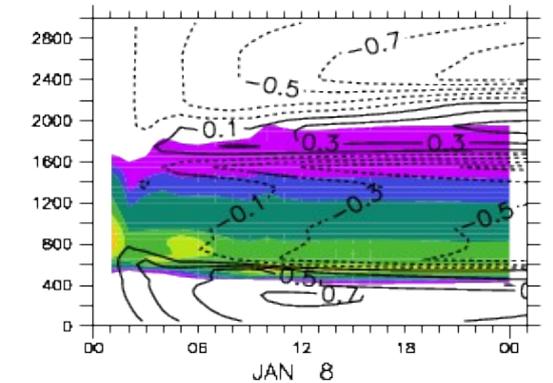
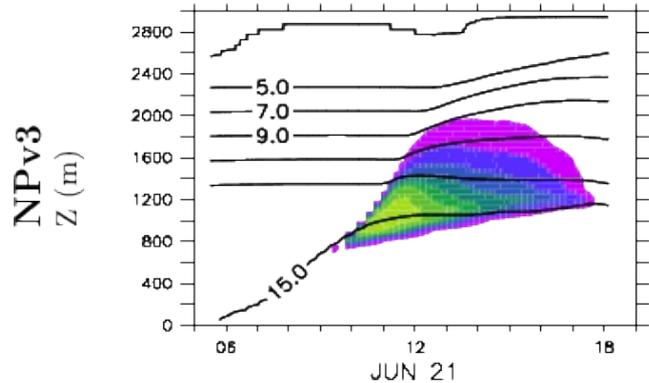
Reference



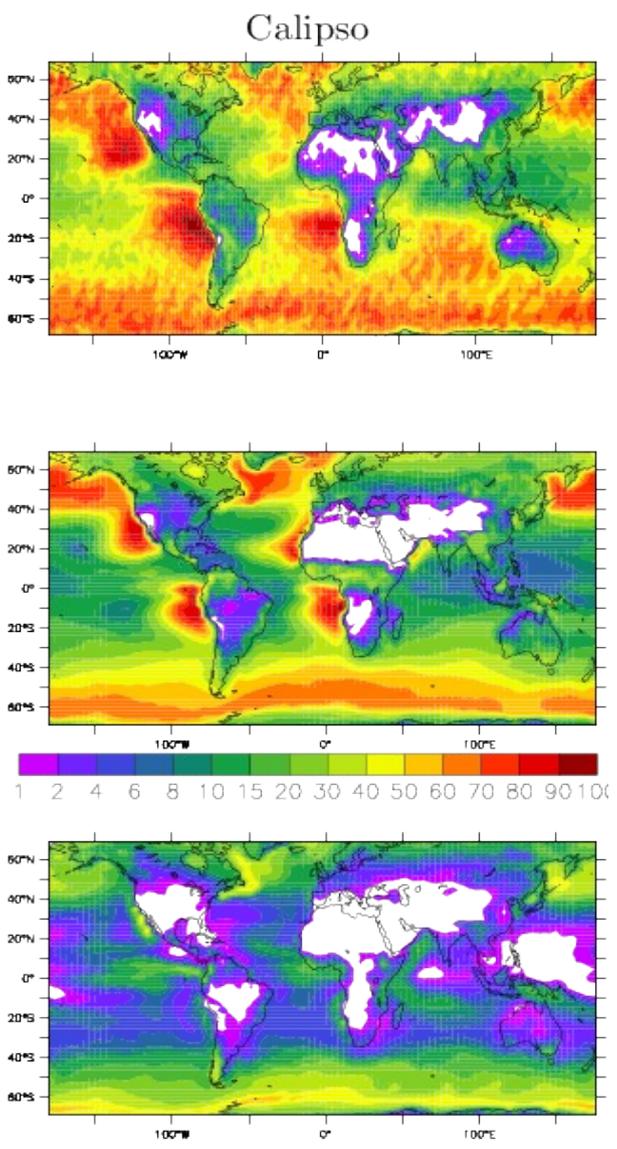
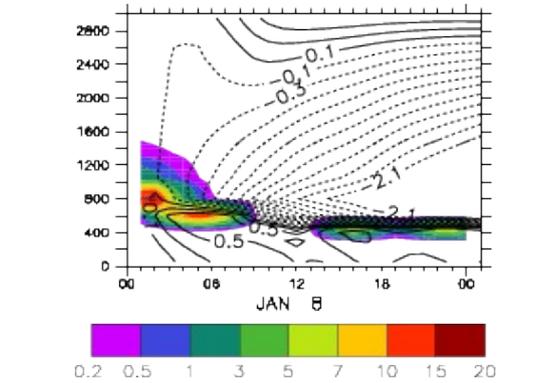
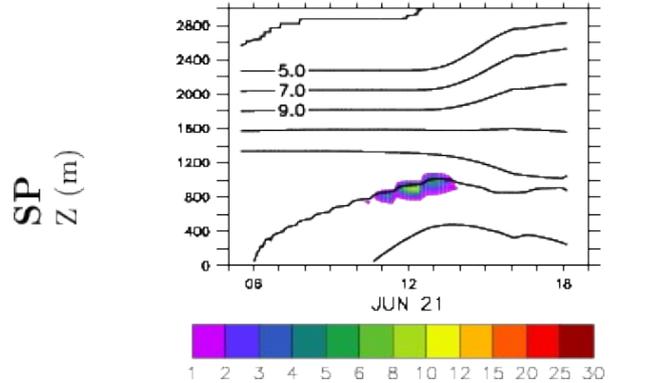
LES



IPSL-CM5B



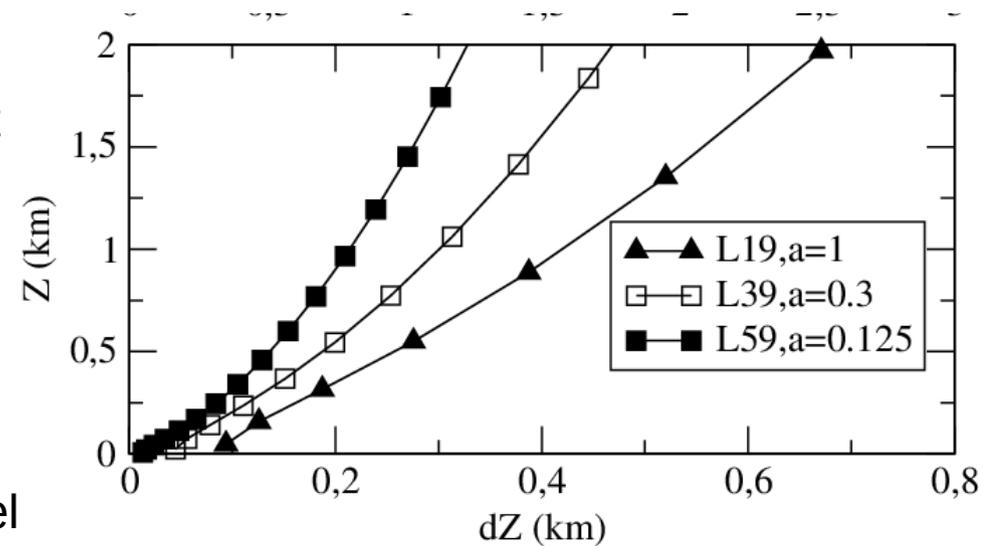
IPSL-CM5A



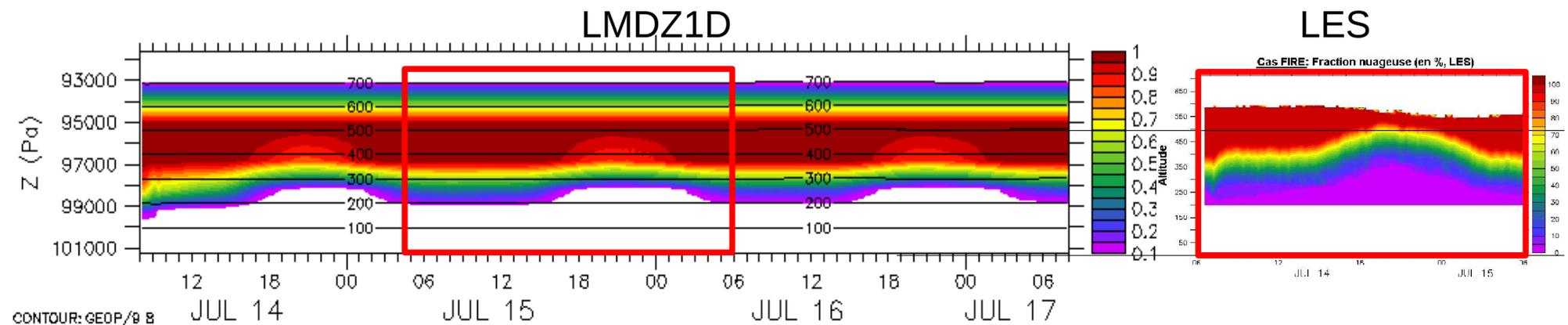
Better representation of low-level clouds in IPSL-CM5B

To come in LMDZ6

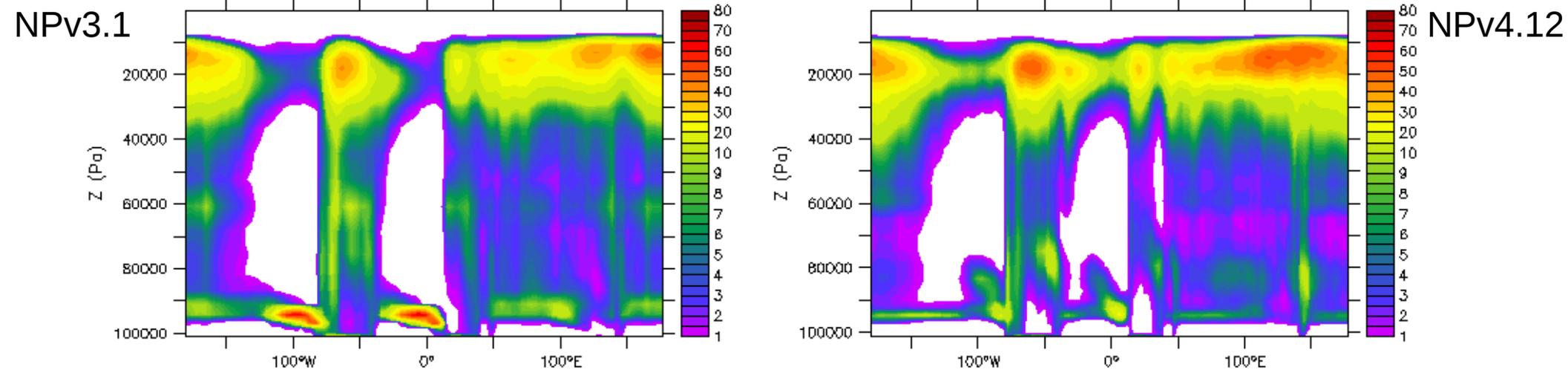
- Increase of vertical resolution with a refinement at low levels
- Activation of the thermal plume everywhere
- Modification of entrainment and detrainment in the thermal plume model to account for cloud-top mixing in stratocumulus (Jam et al., in preparation)

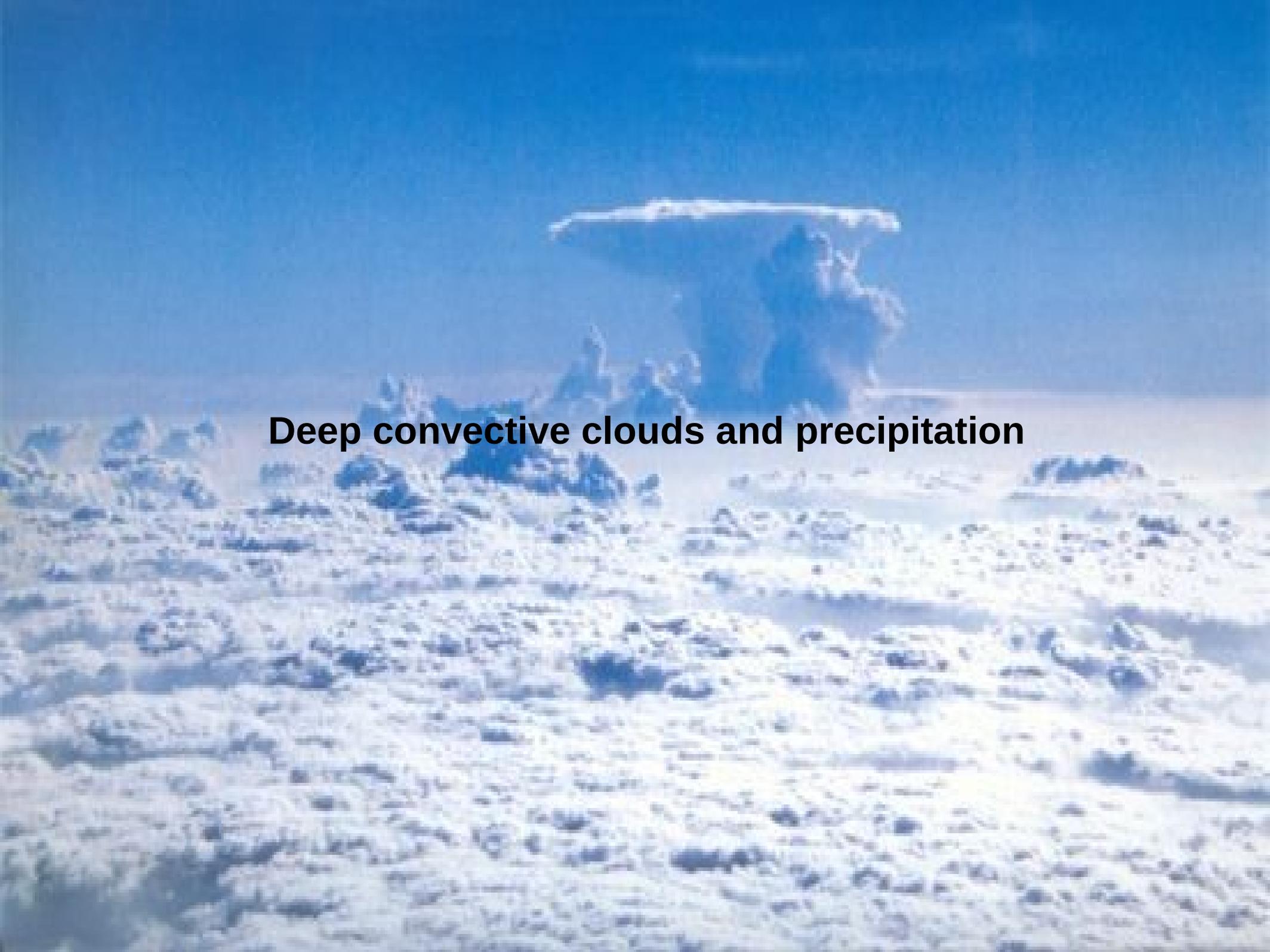


1D FIRE case with the modified thermal plume model



Effect in 3D simulations: Vertical profile of cloud fraction averaged between 5S and 20S





Deep convective clouds and precipitation

Cumulonimbus, updrafts and cold pools

Local convection in semi-arid region: The 10 of July 2006 in Niamey

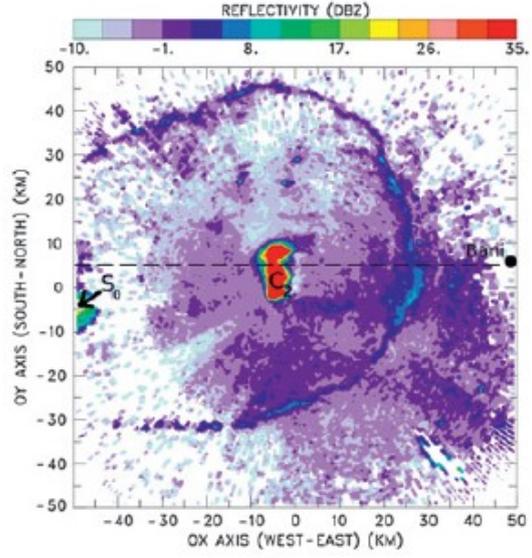
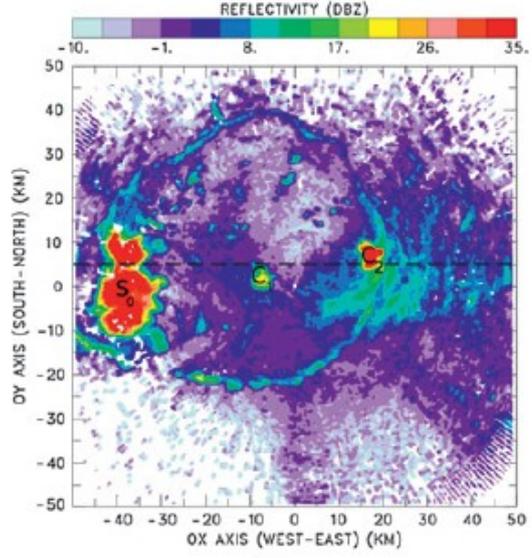
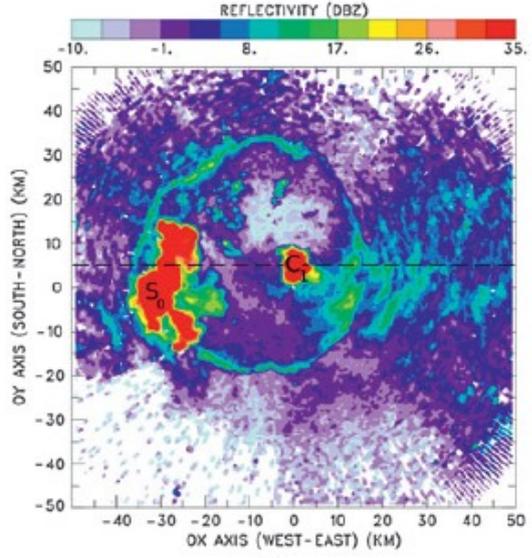
Development of organized structures associated with deep convection

17:20UTC

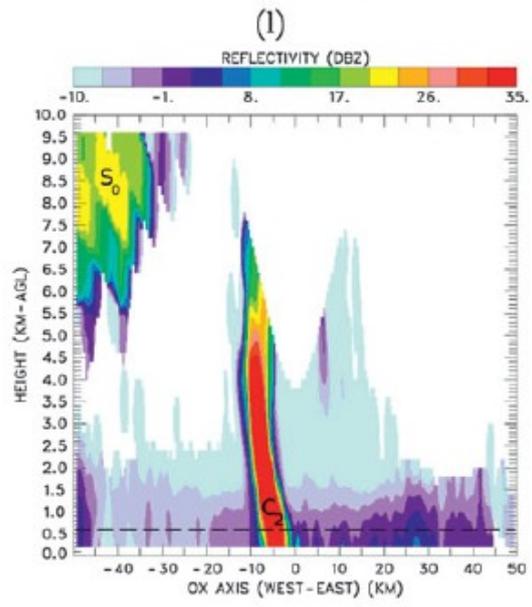
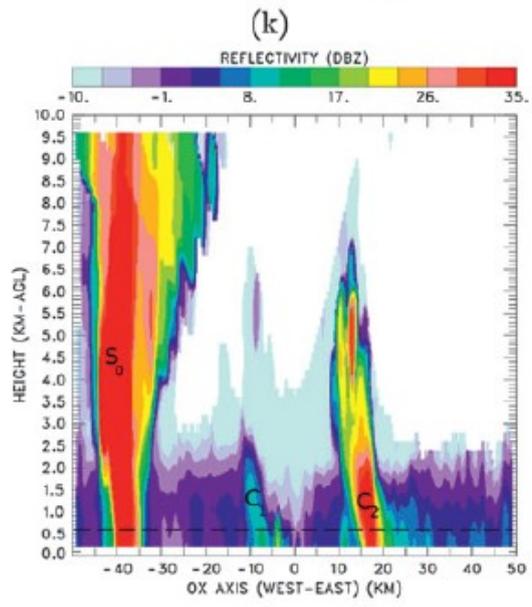
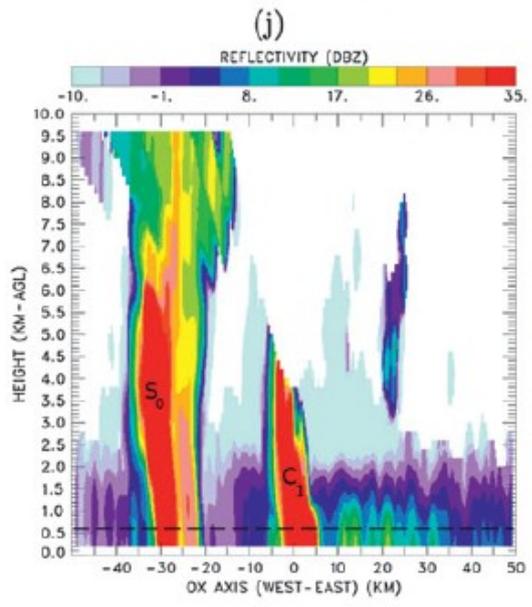
17:40UTC

18:20UTC

Horizontal cross-section at 600m

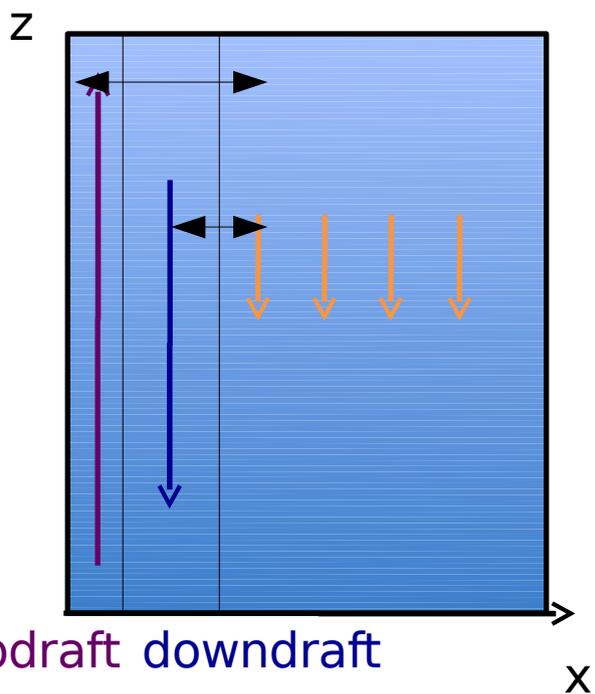


Vertical cross-section 5km north of the RADAR

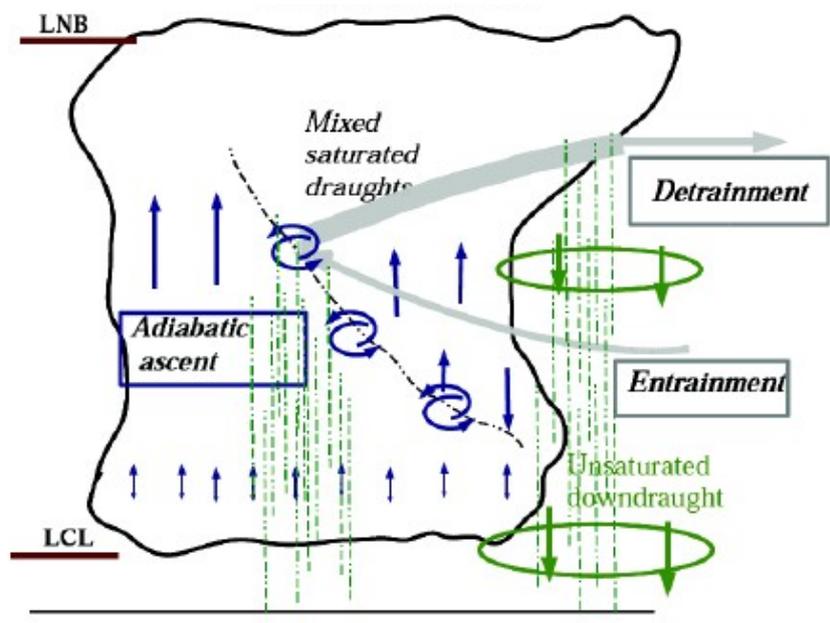


The deep convection scheme

concvl.F



updraft downdraft



Emanuel, 1991

Parameterization of cold pools (LMDZ5B)

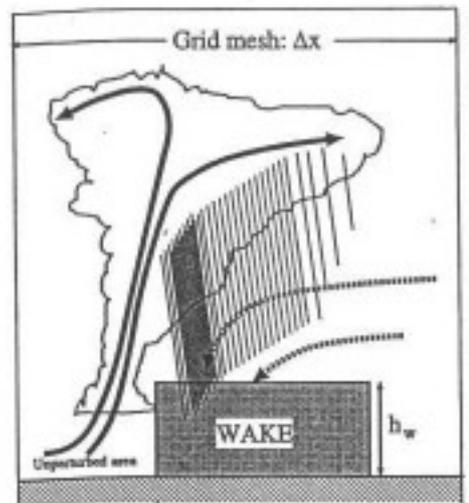
- Triggering function of the deep convection scheme:
Criteria on the convective inhibition

- Convection intensity ("closure"):
Convective intensity related to mean environmental properties (LMDZ5A)
Convective intensity related to sub-cloud processes (LMDZ5B)

- Precipitation efficiency: fraction of condensate that precipitates instead of being detrained

- Updrafts and downdrafts properties: vertical velocity, buoyancy and fractional coverage

- Mixing rates between clouds and environment

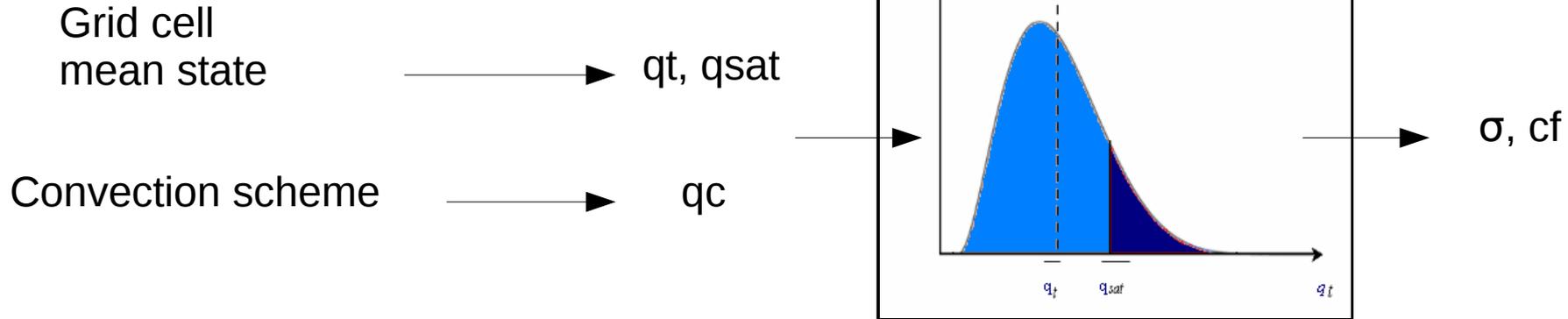


Grandpeix & Lafore, JAS, 2010

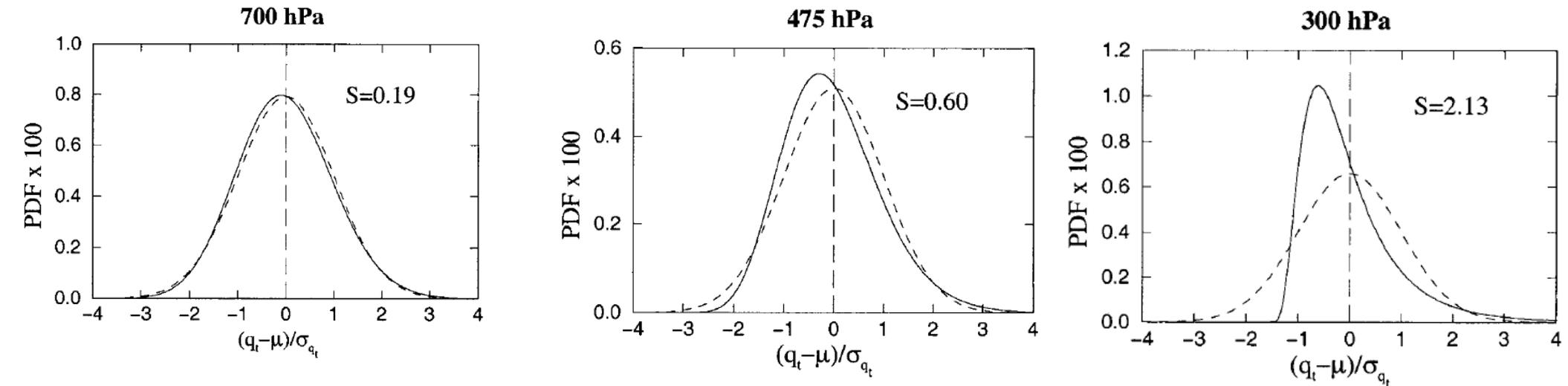
The deep convection cloud scheme

clouds_gno.F

Log-normal distribution of total water q_t



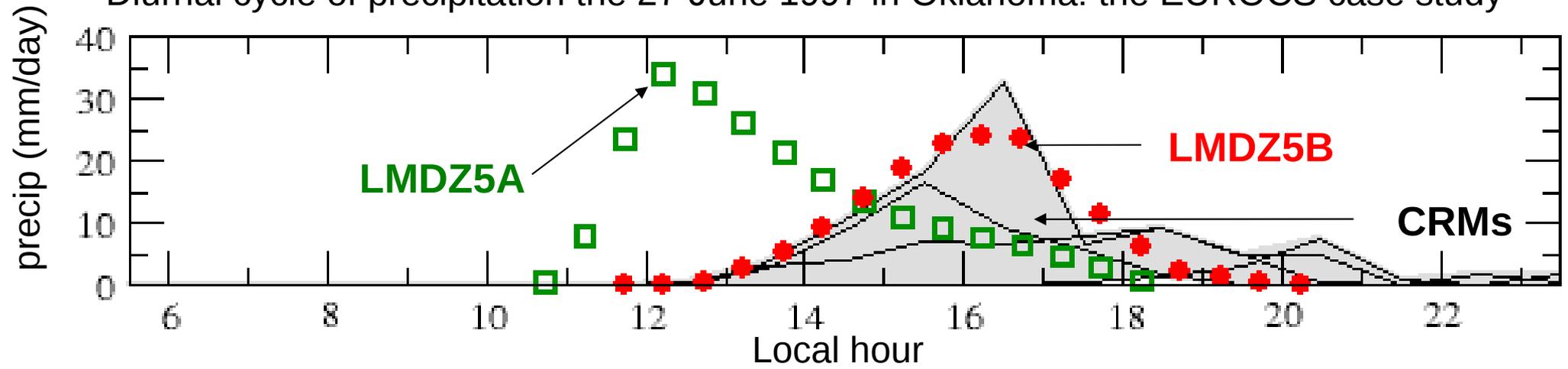
Vertical variation of the PDF on the oceanic case TOGA-COARE
20-27 December 1992



Bony & Emanuel, JAS, 2001

Representation of the diurnal cycle of precipitation over land

Diurnal cycle of precipitation the 27 June 1997 in Oklahoma: the EUROCS case study

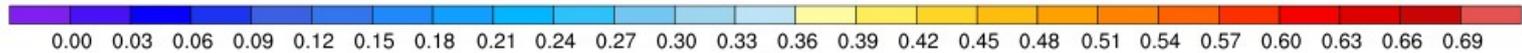
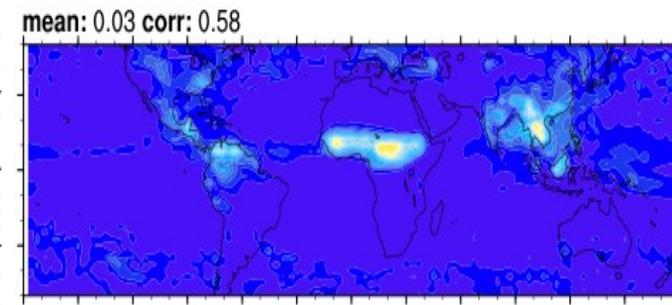
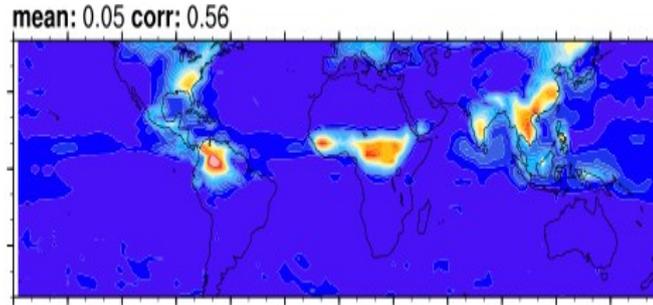
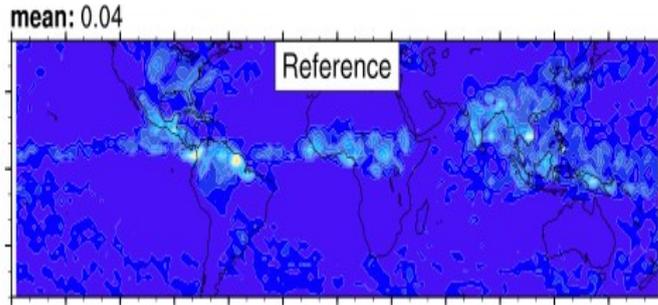


TRMM-3B42

LMDZ5A

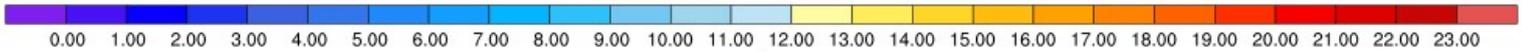
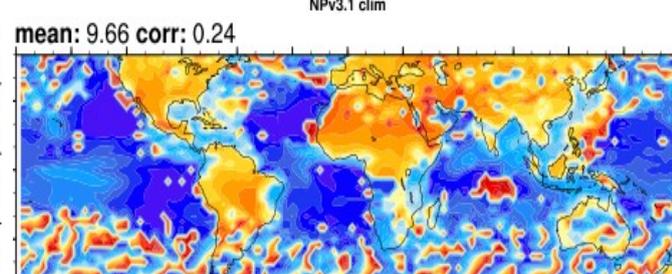
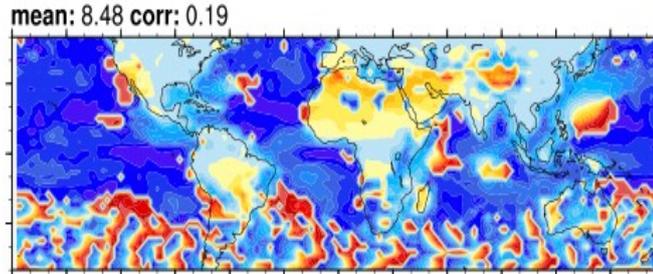
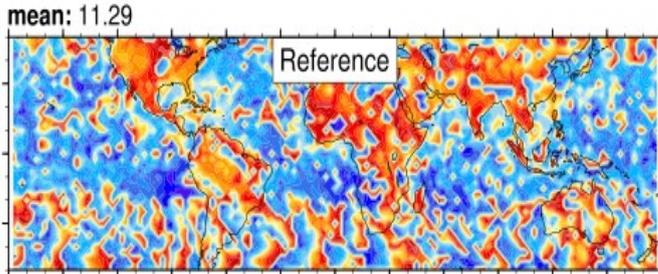
LMDZ5B

Amplitude of maximum rainfall



mm/hour

Local hour of maximum rainfall



mm/hour

To come in LMDZ6

Modification of the **triggering criteria** of the deep convection scheme
(Rochetin et al., JAS, 2014)

- Deterministic approach of deep convection triggering:
 $ALE > |CIN|$

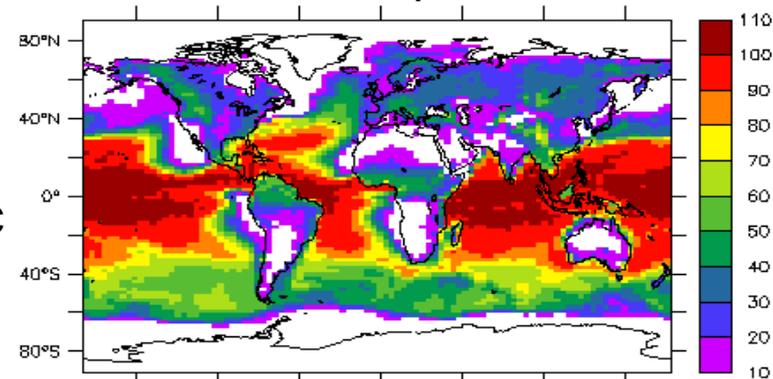
- Probabilistic approach:
What is the probability the grid-cell contains one thermal sufficiently large to trigger convection?

- **Stochastic approach:**

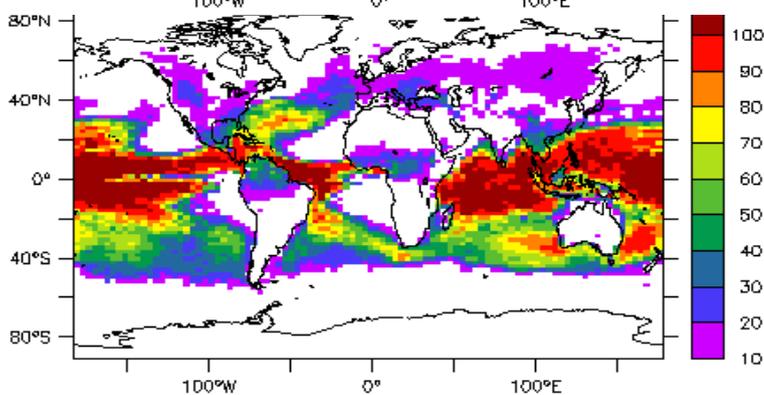
Random number between 0 and 1 to be compared with the triggering probability

Frequency of occurrence of deep convection

deterministic



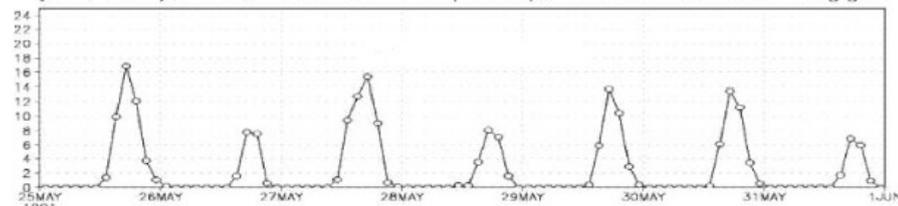
stochastic



Day-to day variability of precipitation

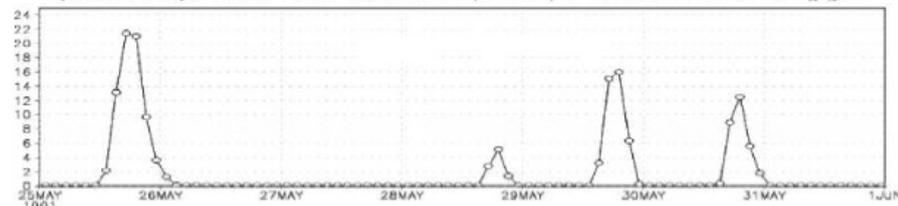
deterministic

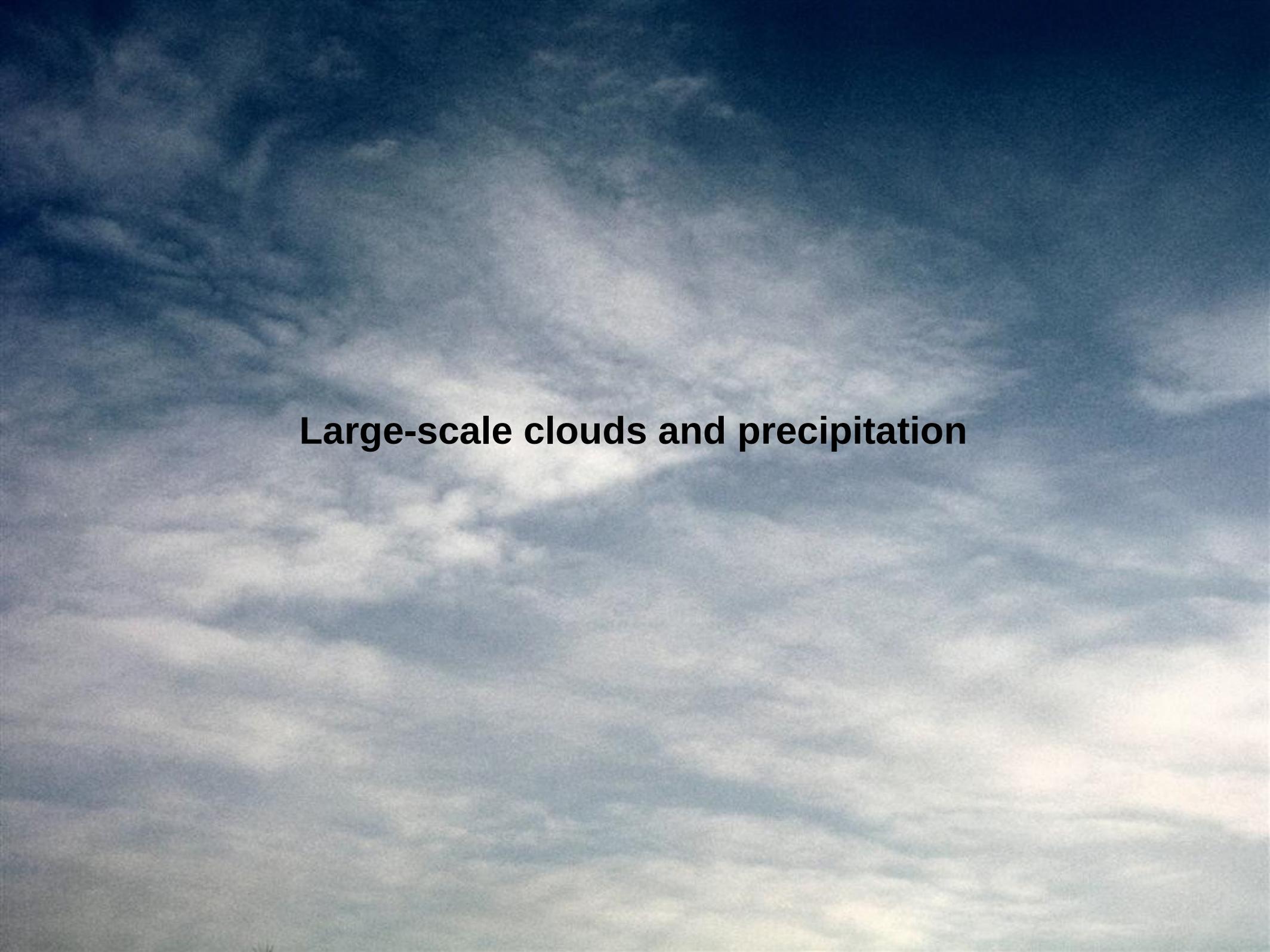
(0E,10N), Jul, Convective precip, deterministic trigger



stochastic

(0E,10N), Jul, Convective precip, stochastic trigger

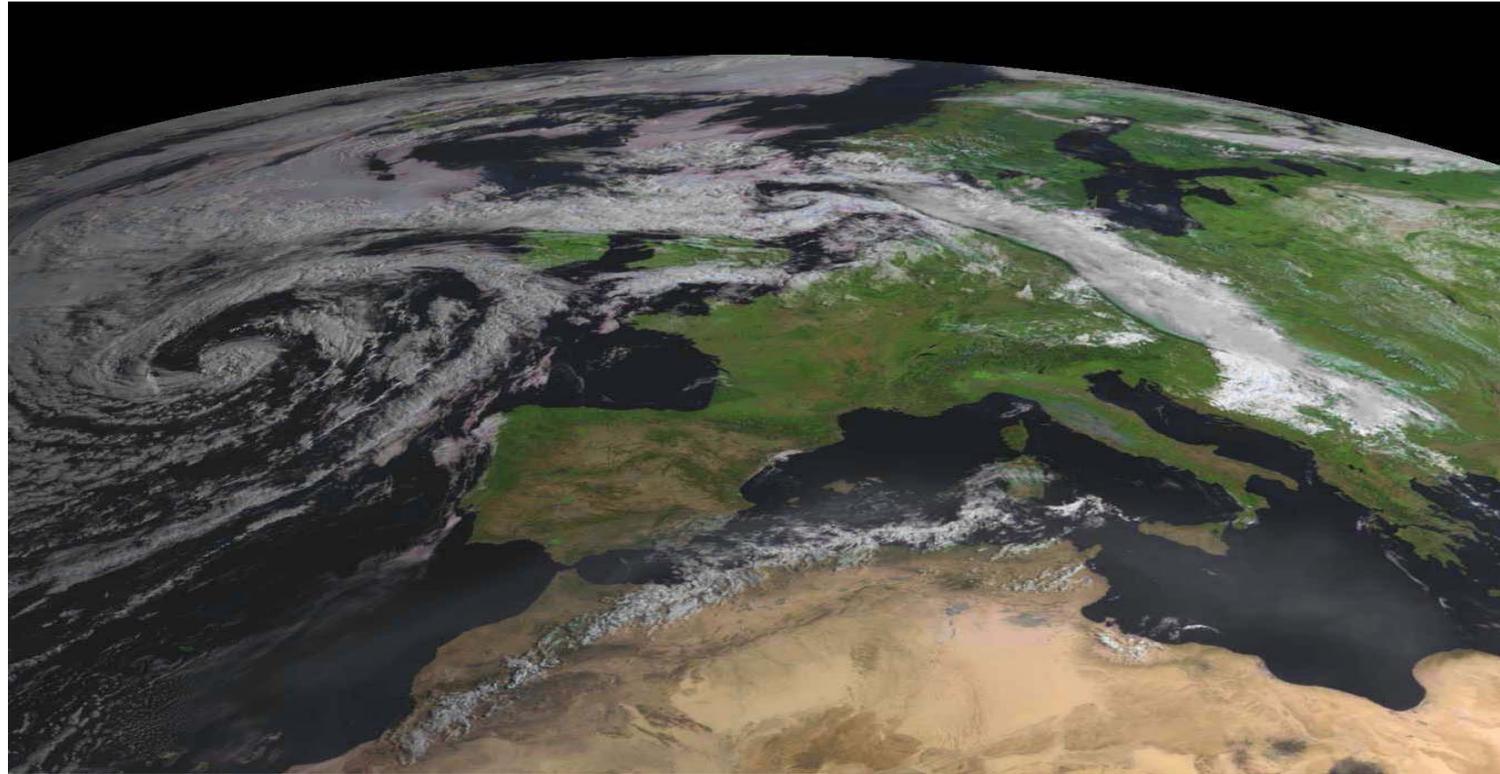




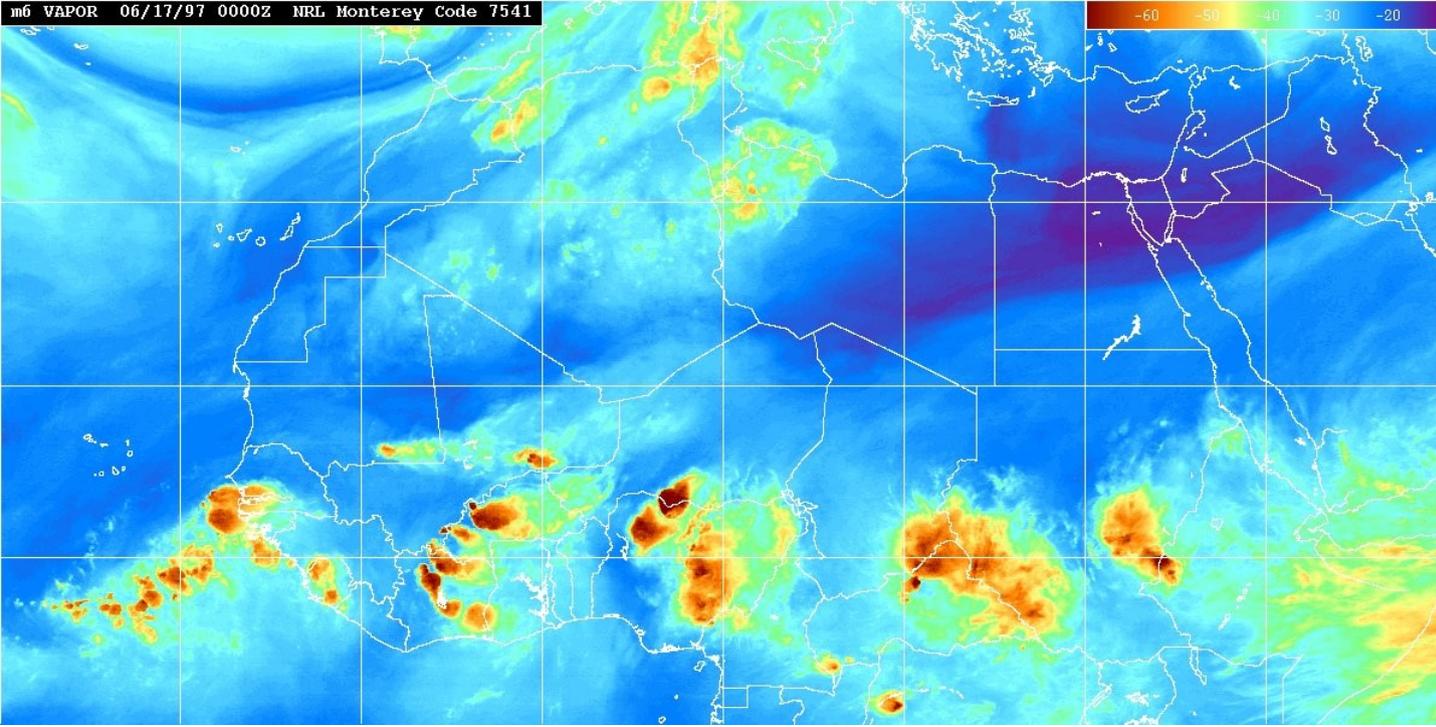
Large-scale clouds and precipitation

Large-scale condensation

Mid-latitude
cyclones



m6 VAPOR 06/17/97 0000Z NRL Monterey Code 7541



Convection organized
in squall lines
in Africa

The cloud scheme

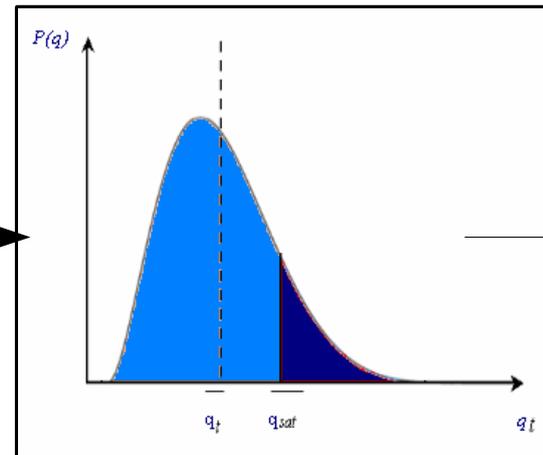
fisrtlp.F90

Log-normal distribution of total water q_t (Bony & Emanuel, JAS, 2001)

Grid-cell
mean state

→ q, q_{sat}

σ/q imposed



$$\alpha_c = \int_{q_{sat}}^{\infty} P(q) dq$$

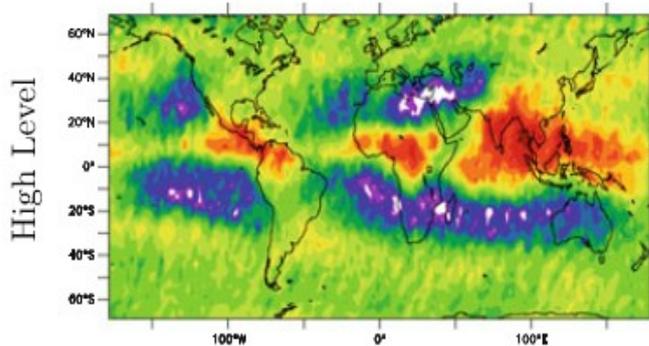
$$q_c = \int_{q_{sat}}^{\infty} (q - q_{sat}) P(q) dq$$

3D simulations

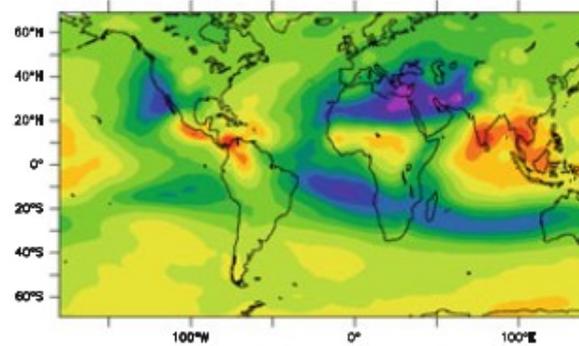
High cloud fraction (%)

Annual mean

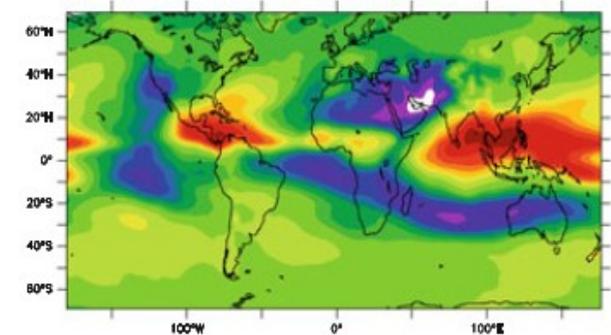
GOCCP



LMDZ5A (SP)



LMDZ5B (NPv3)



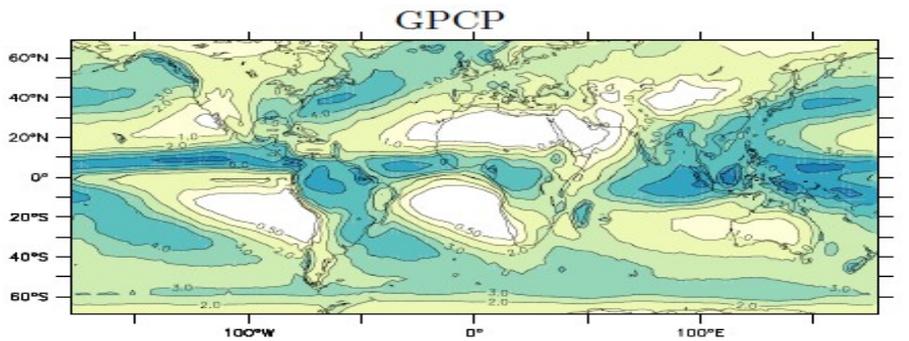
Representation of precipitation mean and variability

Hourdin et al., 2012

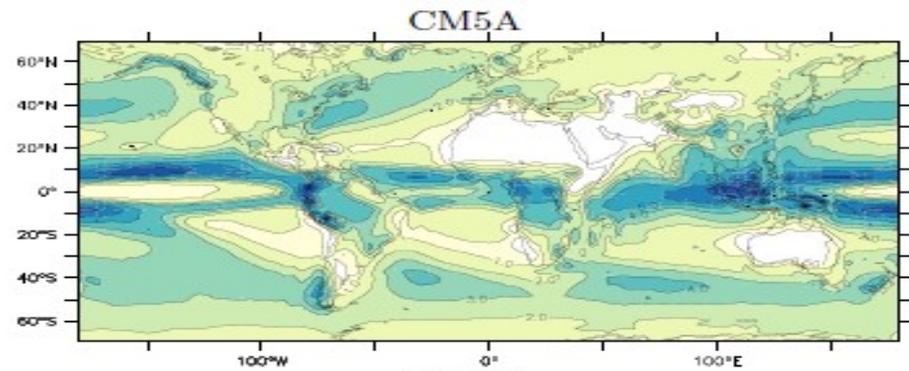
Annual precipitation (mm/day)

Intraseasonal variability of precipitation (mm/day)

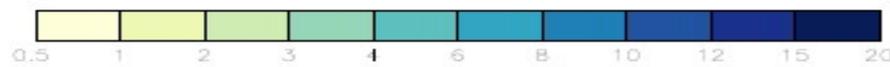
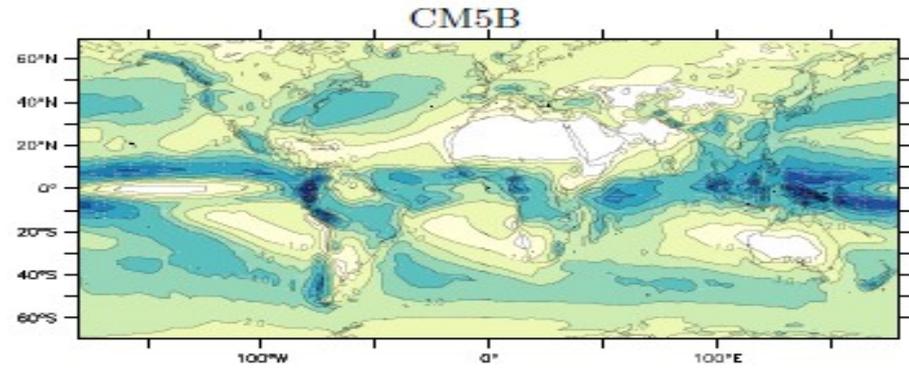
GPCP



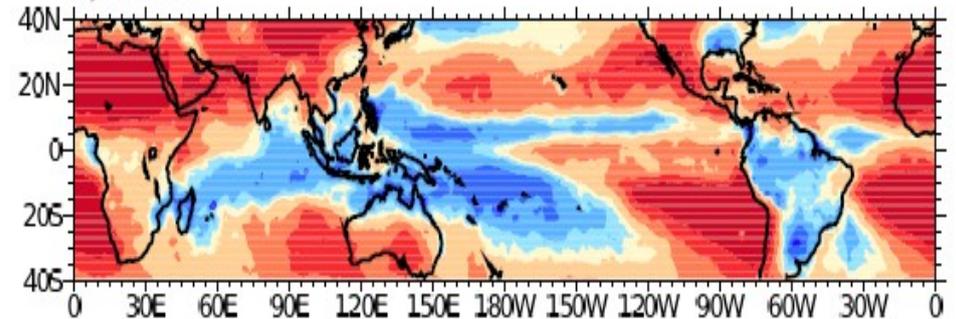
IPSL-CM5A



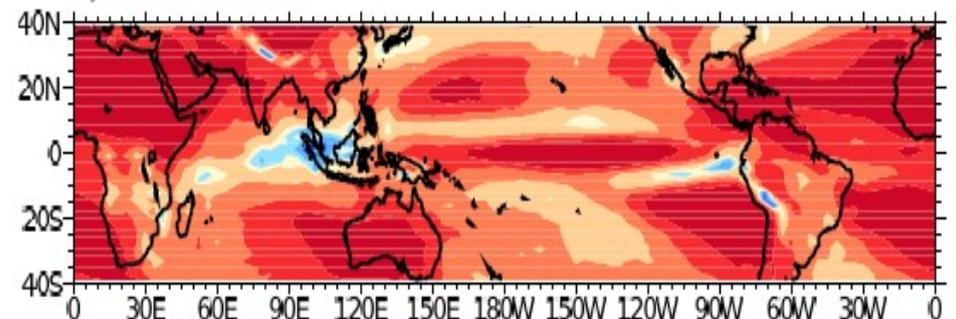
IPSL-CM5B



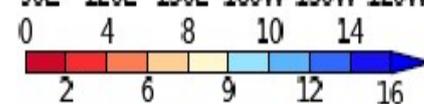
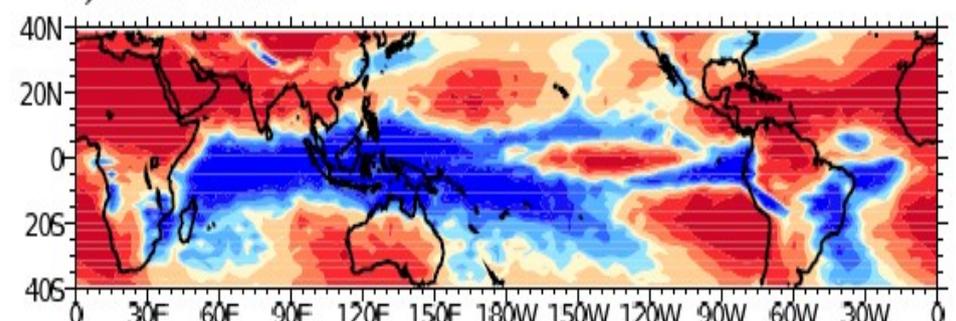
a) GPCP



b) IPSL-CM5A



c) IPSL-CM5B



Reduction of the double-ITCZ problem in IPSL-CM5B

Reinforcement of the intra-seasonal variability of precipitation over ocean in IPSL-CM5B

To come in LMDZ6: The thermodynamical effect of ice

In the deep convection scheme

- Latent heat release associated with freezing in updrafts
Hyp: transformation of liquid to ice between -10 and -40°C
- Absorption of latent heat associated with melting in unsaturated downdrafts
- Introduction of 2 precipitation fluxes:
Liquid and ice
- Different latent heat for evaporation and sublimation

Hyp: Ice melts linearly between 0 and 15°C

Parameter: precipitation efficiency



<http://www.cnrm-game.fr/>

In the large-scale condensation scheme

- Co-existence of liquid and ice between 0 and -40°C .
- Introduction of 2 precipitation fluxes: Liquid and ice
- Different latent heat for evaporation and sublimation

Hyp: Ice melts linearly between 0 and 2°C

Parameters: maximum water content of clouds, autoconversion rate

+ coefficient of evaporation, scaling factor on the falling speed of ice crystals, effective radius of droplets and ice crystals

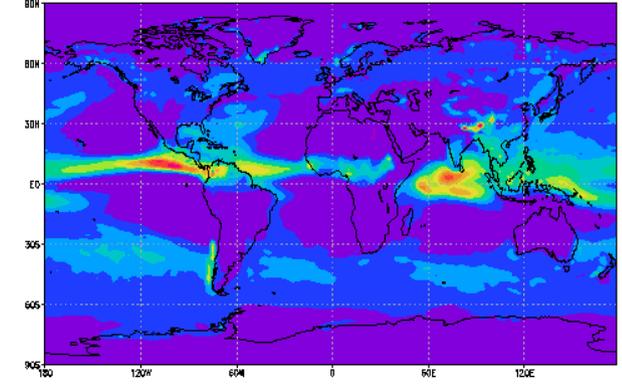
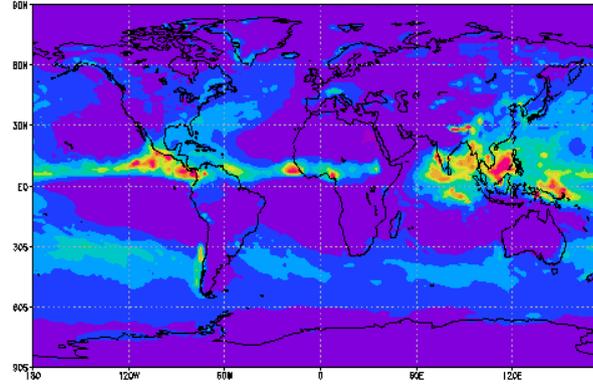
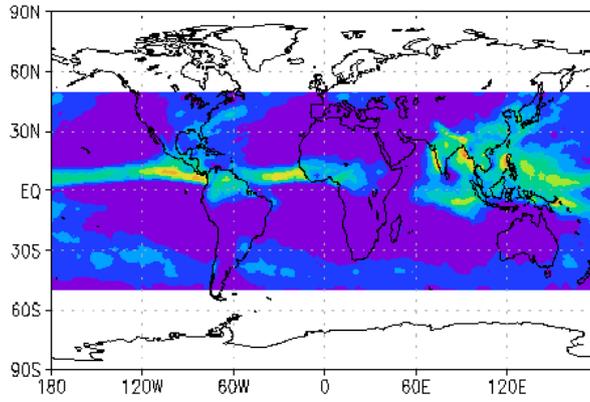
To come in LMDZ6

TRMM

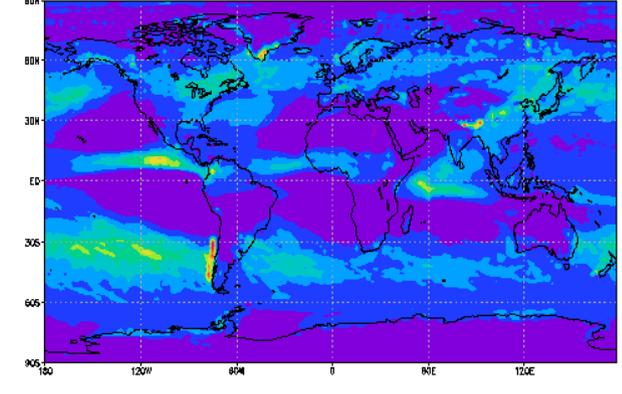
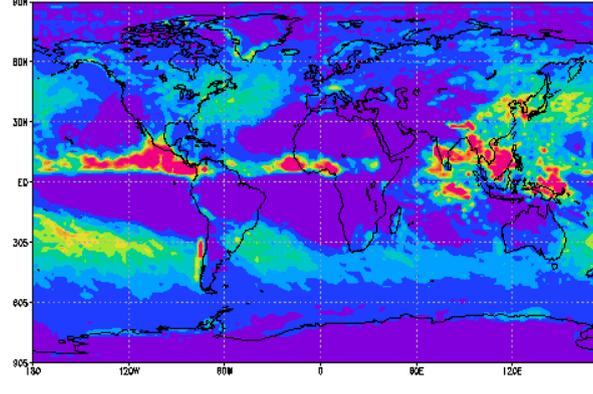
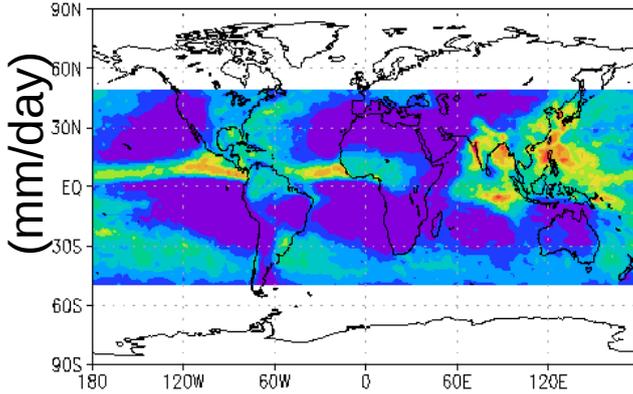
LMDZ5B

NPv5.17c

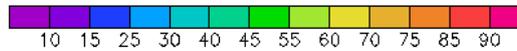
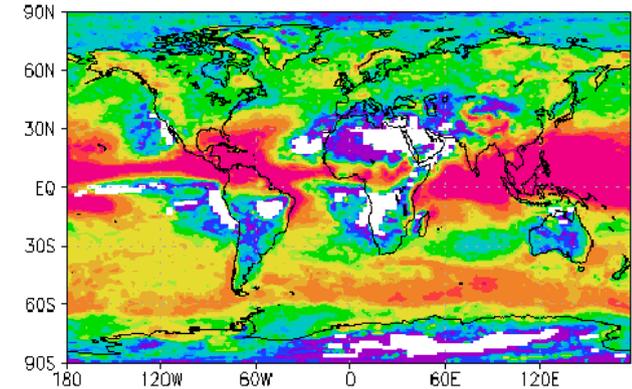
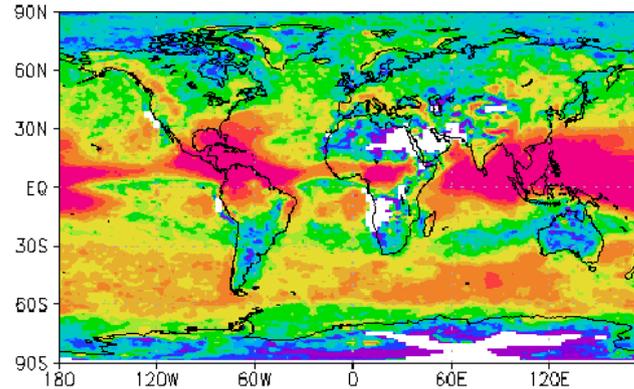
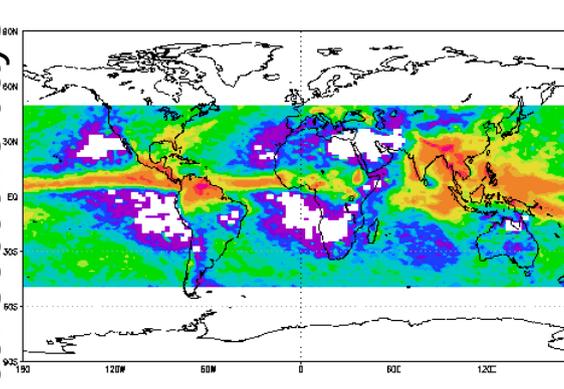
Mean (mm/day)



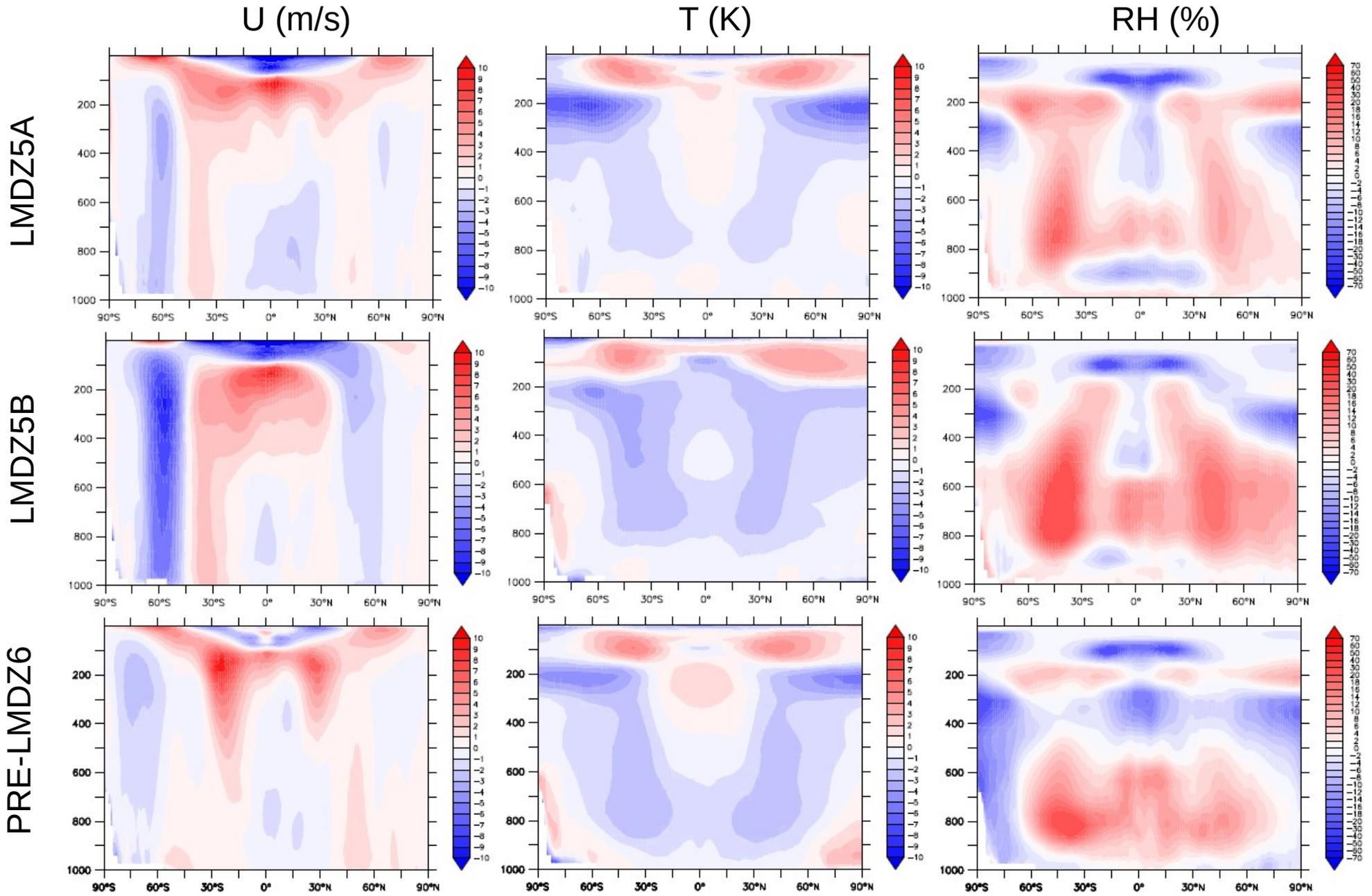
Standard deviation (mm/day)



Probability of 2 consecutive days (%)



What about mean biases?

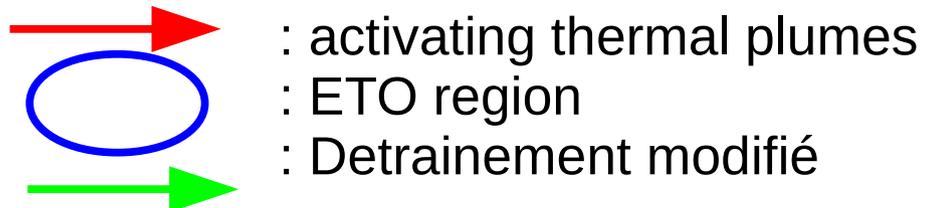


Mean biases amplified in LMDZ5B compared to LMDZ5A
Diminution of mean biases in PRE-LMDZ6

What about mean biases?

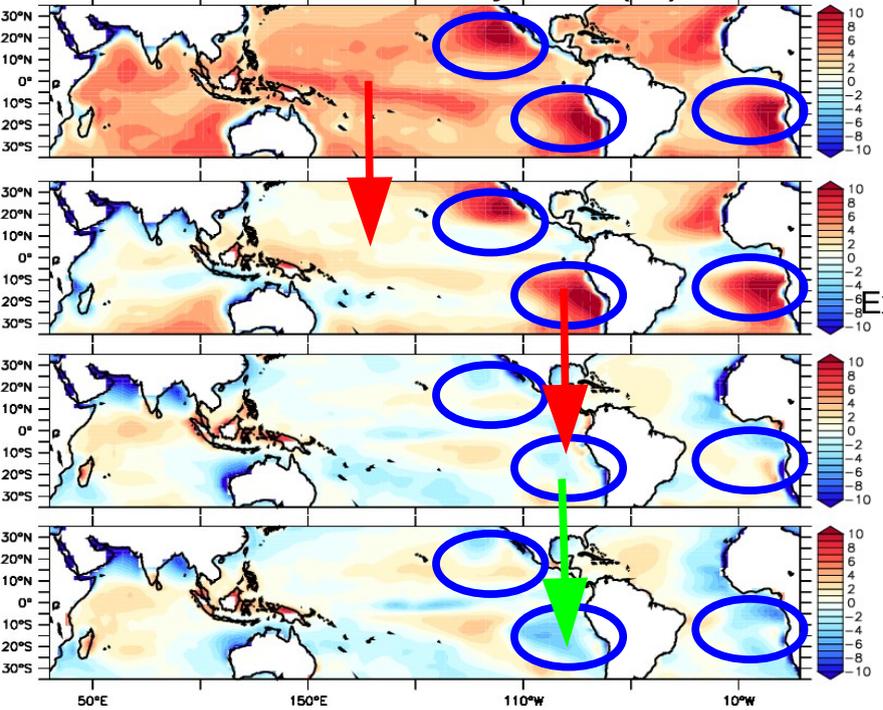
Results from atmospheric simulations forced by climatic sea surface temperature

Impact of the thermal plume model on the near surface humidity bias

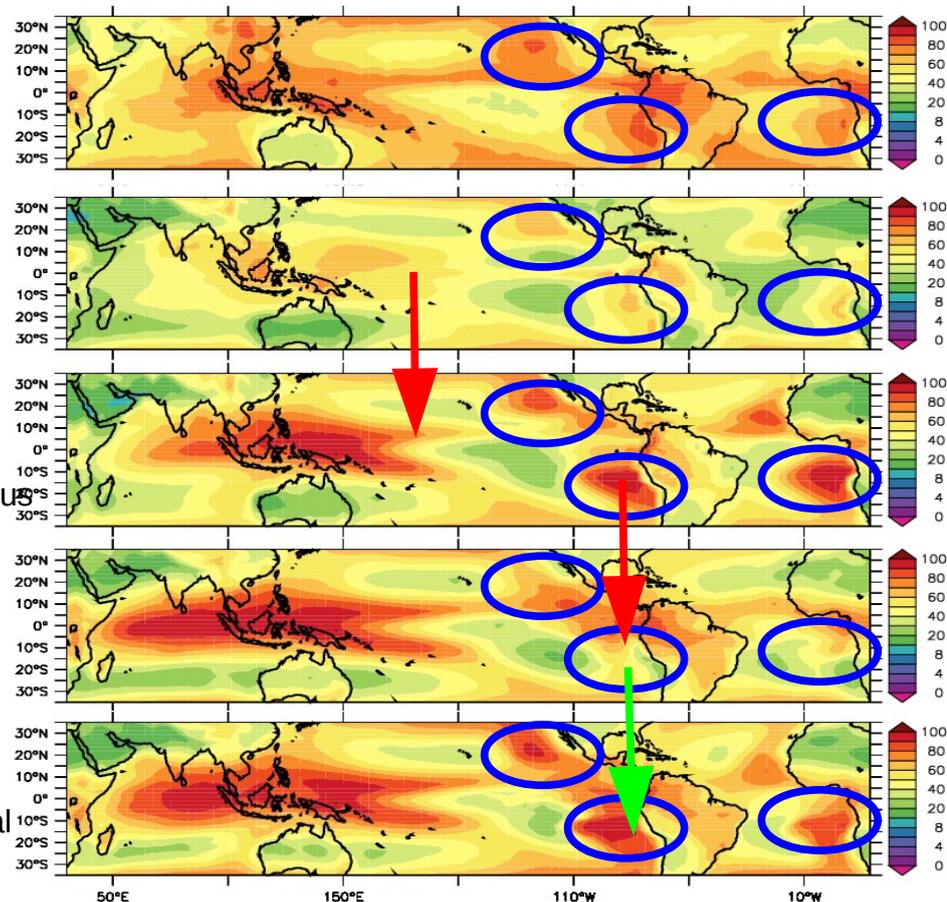


Observations Calipso GOCCP
Da Silva

Relative humidity bias (%)



Total cloud cover (%)



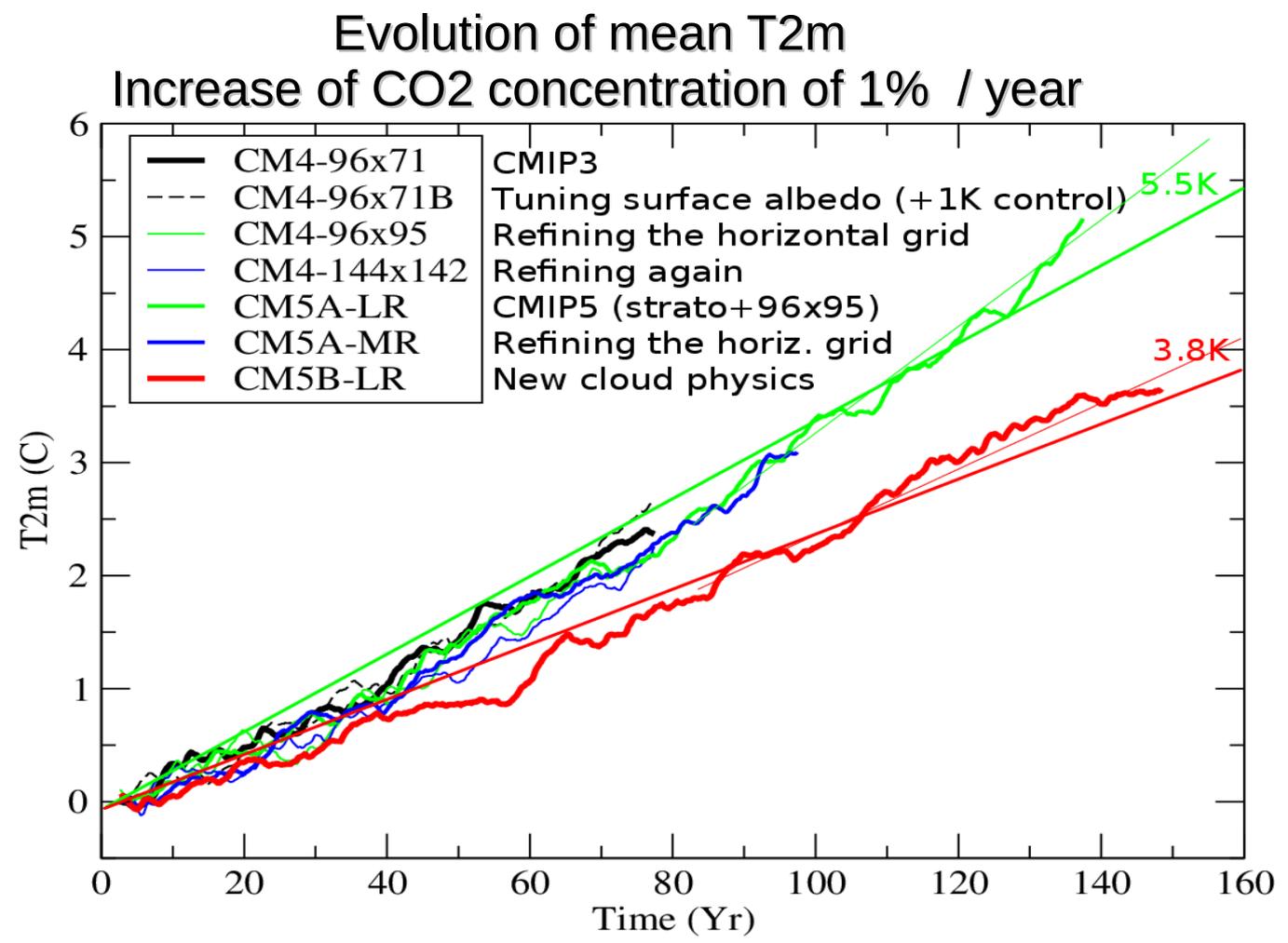
To come in LMDZ6

Reduction of the relative humidity bias in the Eastern part of Tropical Ocean
Reduction of corresponding latent heat flux bias in forced mode and SST bias in coupled mode

What about thermohaline circulation and climate sensitivity?

- ✓ Dramatic drop of the thermohaline circulation in IPSL-CM5B
- ✓ Lower sensitivity in IPSL-CM5B than in IPSL-CM5A

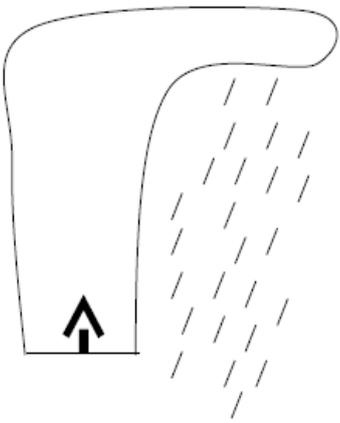
Sensitivity to resolution and physical package



Hourdin et al., 2012

To come in LMDZ6?

The different physical packages of LMDZ

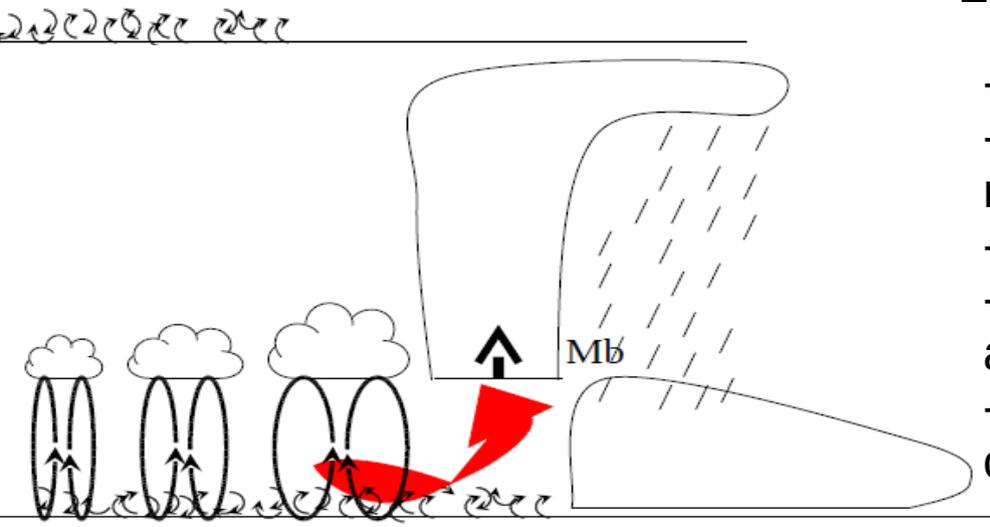


LMDZ5A

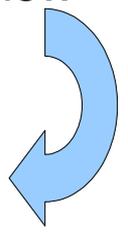
- Diffusion scheme (Louis, 1979)
- Deep convection (Emanuel, 1991)
- Cloud scheme (Bony et Emanuel, 2001)



LMDZ5B



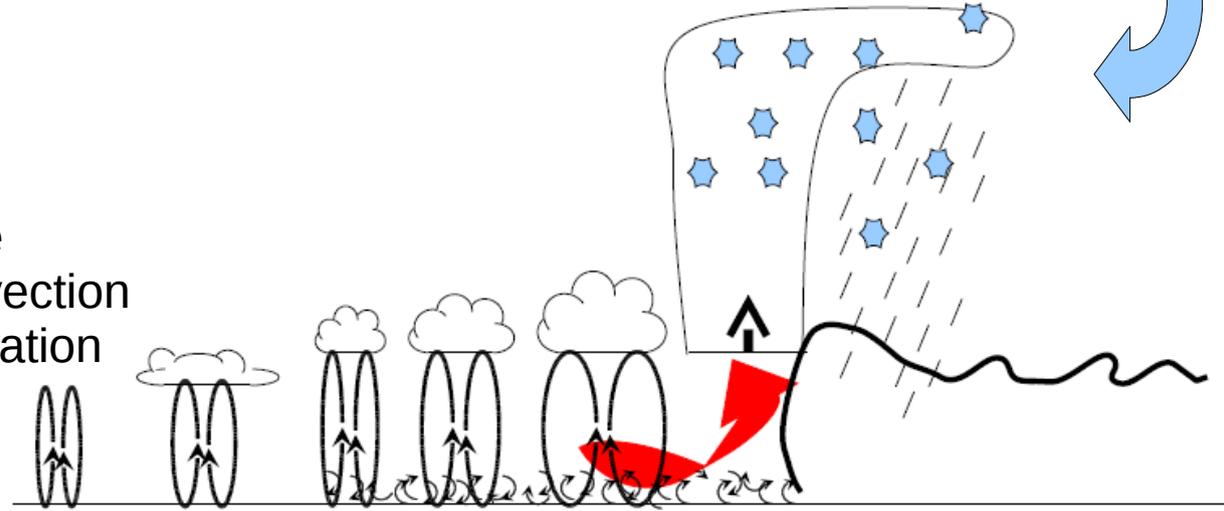
- Diffusion scheme (Yamada, 1983)
- Thermal plume model in shallow cumulus regions (Rio et al., 2010)
- Cold pool (Grandpeix et Lafore, 2010)
- Deep convection controlled by thermals and wakes (Rio et al., 2012)
- Bi-gaussian cloud scheme for shallow convection (Jam et al., 2013)



PRE-LMDZ6

LMDZ5B

- + Thermal plume model everywhere
- + Stochastic triggering of deep convection
- + Different convective mixing formulation
- + Thermodynamical effect of ice



LMDZ5A (AR4_physiq.def)

iflag_pbl=1
iflag_thermals=0
iflag_thermals_ed=0
iflag_coupl=0

iflag_con=30
iflag_clos=1
iflag_wake=0

iflag_mix=0
iflag_clw=1
epmax=0.999

iflag_cldcon=3
iflag_ratqs=0
ratqsbas=0.005
ratyqshaut=0.33

cld_lc_lsc=4.16e-4
cld_lc_con=4.16e-4
ffallv_lsc=0.5
ffallv_con=0.5
coef_eva=2e-5

Boundary-layer

Diffusion
Thermals
Mixing rates in thermals
Coupling with deep convection

Convection

Emanuel old/new
Closure CAPE/ALP
Cold pools

PDF for mixing
Computation of condensate
Efficiency of precipitation

Clouds

Cloud scheme
Profile of σ/qt
 σ/qt min
 σ/qt max

Threshold cloudy water LS
Threshold cloudy water CV
Ice crystals fall speed LS
Ice crystals fall speed CV
Coefficient of evaporation

LMDZ5B (NPv3.1_physiq.def)

iflag_pbl=8
iflag_thermals=15
iflag_thermals_ed=10
iflag_coupl=5

iflag_con=3
iflag_clos=2
iflag_wake=1

iflag_mix=1
iflag_clw=0
epmax=0.997

iflag_cldcon=6
iflag_ratqs=2
ratqsbas=0.002
ratqs_haut=0.25

cld_lc_lsc=6e-4
cld_lc_con=6e-4
ffallv_lsc=1.35
ffallv_con=1.35
coef_eva=1e-4

LMDZ5A (AR4_physiq.def)

iflag_pbl=1
iflag_thermals=0
iflag_thermals_ed=0
iflag_coupl=0

iflag_con=30
iflag_clos=1
iflag_wake=0

iflag_mix=0
iflag_clw=1
epmax=0.999

iflag_cldcon=3
iflag_ratqs=0
ratqsbas=0.005
ratyqshaut=0.33

cld_lc_lsc=4.16e-4
cld_lc_con=4.16e-4
ffallv_lsc=0.5
ffallv_con=0.5
coef_eva=2e-5

Boundary-layer

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Threshold cloudy water CV
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Ice crystals fall speed CV
Coefficient of evaporation

LMDZ5B (Npv3.1_physiq.def)

(NPv5.17h_physiq.def)

UNDER DEVELOPMENT!!

iflag_pbl=8 (11)
iflag_thermals=15 (18)
iflag_thermals_ed=10 (8)
iflag_coupl=5

iflag_con=3
iflag_clos=2
iflag_wake=1
iflag_trig_bl=1

iflag_mix=1 (0)
iflag_clw=0
Epmax=0.997 (0.998)

iflag_cldcon=6
iflag_ratqs=2 (4)
ratqsbas=0.002
ratqs_haut=0.25 (0.312)
iflag_t_glance=1
iflag_ice_thermo=1

cld_lc_lsc=6e-4 (3e-4)
cld_lc_con=6e-4 (3e-4)
ffallv_lsc=1.35 (0.665)
ffallv_con=1.35 (0.665)
coef_eva=1e-4 (2e-5)